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J. A. JOHANNESSEN & H. G. GADE
A case study of the variations in the upper ocean at
Ocean Weather Ship Mike (66°N, 2°E) in the Norwegian Sea
165

R. FJØRTOFT, D. BJØRGE & K. A. IDEN
Probability models in precipitation climate studies
177

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A case study of the variations in the upper ocean at Ocean Weather Ship Mike (66° N, 2° E) in the Norwegian Sea

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This paper investigates how short period variations in the upper ocean structure relates to meteorological changes. The one-dimensional, mixed layer model of Denman has been tested since it requires only routine ocean weather ship observations and no current measurements as input. Agreement between observations and model simulations is obtained on a weekly time scale. Rapid fluctuations (daily) in the depth and temperature of the mixed layer are observed but not successfully modelled. It is suggested that these rapid fluctuations are associated with frontal dynamics such as formations of meanders and eddies.

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1. INTRODUCTION

Since 1948 regular oceanographic observations have been carried out at Ocean Weather Ship (OWS) Mike (66° N, 2° E) in the Norwegian Sea. During these years an appreciable amount of data has been collected. These data have formed studies of the deep water and the Atlantic water in the Norwegian Sea such as by Mosby (1959, 1979). Furthermore, studies of the seasonal variations in the upper layer at OWS Mike from 1948–1958 have been carried out by Helland (1963), while Bøyum (1966) has studied the annual variations in the energy exchange across the air–sea interface over the same period. Recently Gammelsrød & Holm (1983) have looked for possible influence of advection on the upper ocean struc-

ture as opposed to atmospheric forcing over a period of nearly 30 years.

Few studies, however, exist which emphasize synoptic variations in the upper ocean structure in this region. This paper briefly describes the regional background in Sect. 2 and presents mean seasonal variations in the upper ocean structure based on 9 years of data from 1967 to 1975 in Sect. 3. The development of the mixed layer depth and temperature as derived from bathythermograms is examined during the 17 day period 5–22 March 1967 in Sect. 4. This period was selected because of frequent passages of cyclonic storms in the region, and also because the atmosphere gained heat from the sea. The observed behaviour of mixed layer is then simulated by using Denman's (1973) one-dimensional model in Sect. 5.

This model only requires standard marine meteorological parameters as input data. The model prediction of the mixed layer behaviour is compared with observed layer depths and temperatures, and areas of discrepancies are discussed in Sect. 6.

2. REGIONAL BACKGROUND

OWS Mike is located at 66°N , 2°E in the Norwegian Sea. This is in the western boundary region of the Norwegian Atlantic Current (N.A.C.) (Fig. 1). Colder water of Arctic origin is found to the west. The Atlantic water is confined to the upper 400 m at OWS Mike. According to Mosby (1970), the mean speed of the N.A.C. at OWS Mike is 0.03 m/s in a northward direction. Mosby also estimated the average temperature gradient in the Norwegian Sea to be 0.4° per degree latitude. The temperature change to be expected from advective effects would therefore be about 0.10°C in 10 days.

However, both Sælen (1963) and Mosby

(1970) revealed from drawings of depth contours of isothermal surfaces that eddy-like features were occasionally present in the Atlantic current. Such features were first detected in the Norwegian Sea by Helland-Hansen & Nansen (1909) and given the name 'puzzling waves'. The scale of these waves appeared to be in the order of 50 km. Sælen (1963) observed that they could be stationary for several days or moving slowly or rapidly in a north easterly direction. Such eddy-like features within the Atlantic current can be associated with strong horizontal and vertical temperature gradients that may contribute significantly to synoptic variations in the temperature and depth of the mixed layer. This is further discussed in Sect. 6.

3. MEAN SEASONAL VARIATIONS

The mean seasonal variations in the temperature of the upper ocean at selected depths of 0, 50, 150, 200 and 250 m as derived from 9 years of hydrocast data from 1967 to

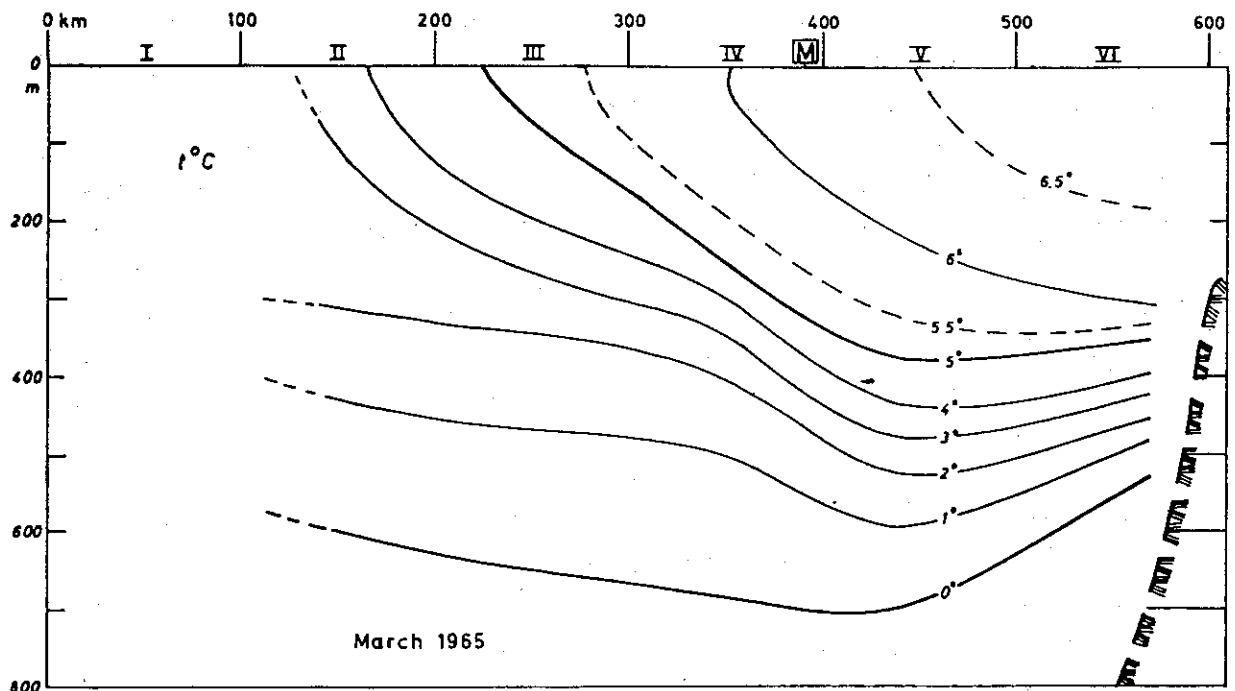


Fig. 1. Average temperature section through OWS Mike (after Mosby 1970).

1975 are presented in Fig. 2a. Based on these curves the active layer at OWS Mike is determined to have a lower limit at 200 m. Below this depth the annual temperature cycle is practically absent. The temperature at this depth varies no more than 1.5°C during the mean year, while the annual fluctuation of the sea surface temperature reaches nearly 6°C. The density variations in this active layer are almost entirely determined by temperature variations. For a more detailed discussion of the seasonal and annual variations of temperature, salinity and density in the upper layer (200 m), the reader is referred to Helland (1963) and Gammelsrød & Holm (1983).

The seasonal variations of the heat flux (q_s) across the ocean-atmosphere interface are presented in Fig. 2b. The heat flux (q_s) is the balance of solar radiation, short-wave reflection from the surface, long-wave radiation from the sky, back radiation and transfer of sensible and latent heat across the sea surface. The period of summer heating, when $q_s > 0$, lasts from May to August with an associated rapid increase in sea surface temperature. From the beginning of September to February the surface temperature decreases, corresponding to the seasonal cooling when $q_s < 0$.

During the ocean heating process the temperature difference between the sea surface

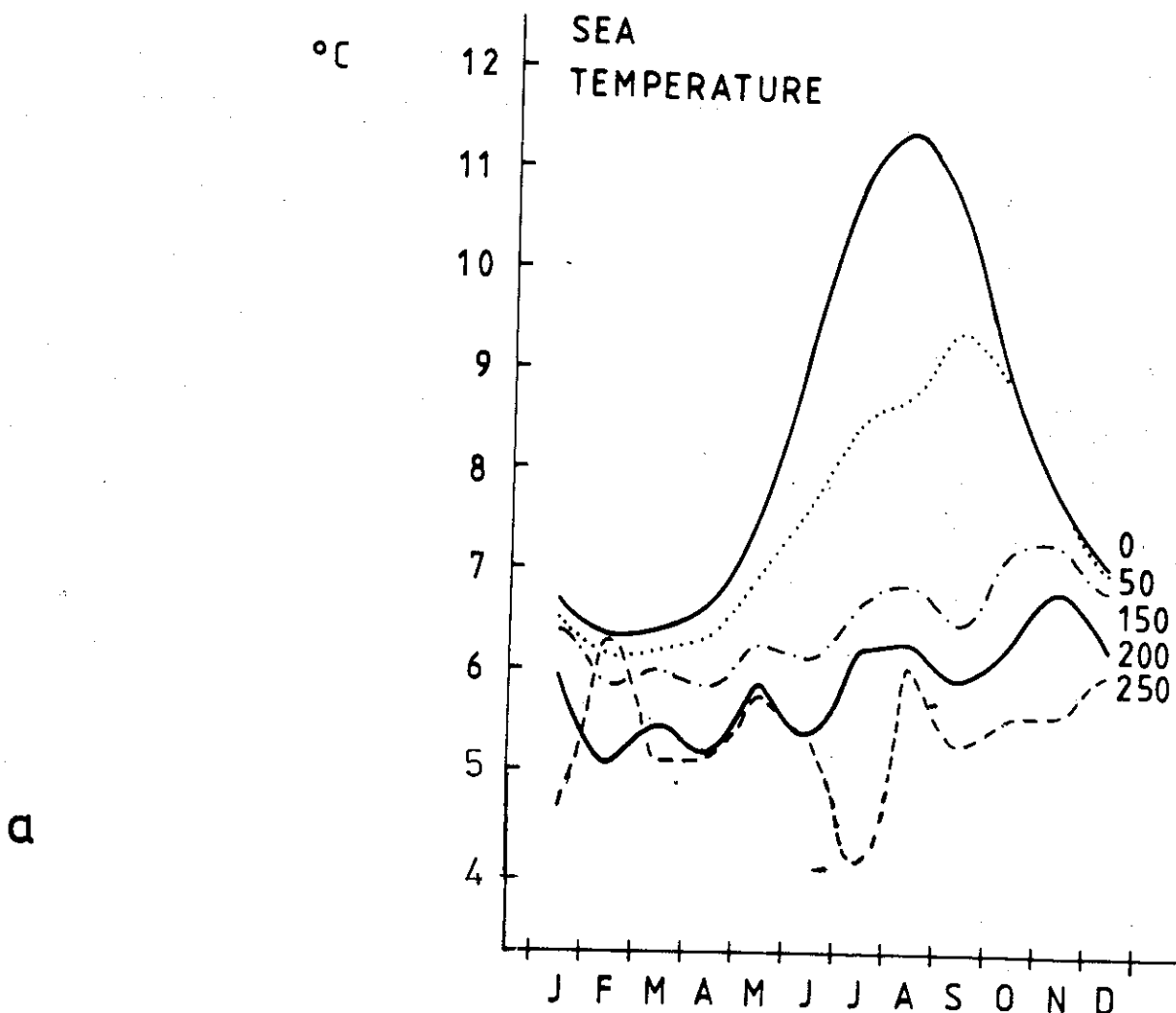


Fig. 2a. Mean seasonal temperature variations with depth for the upper 250 m.

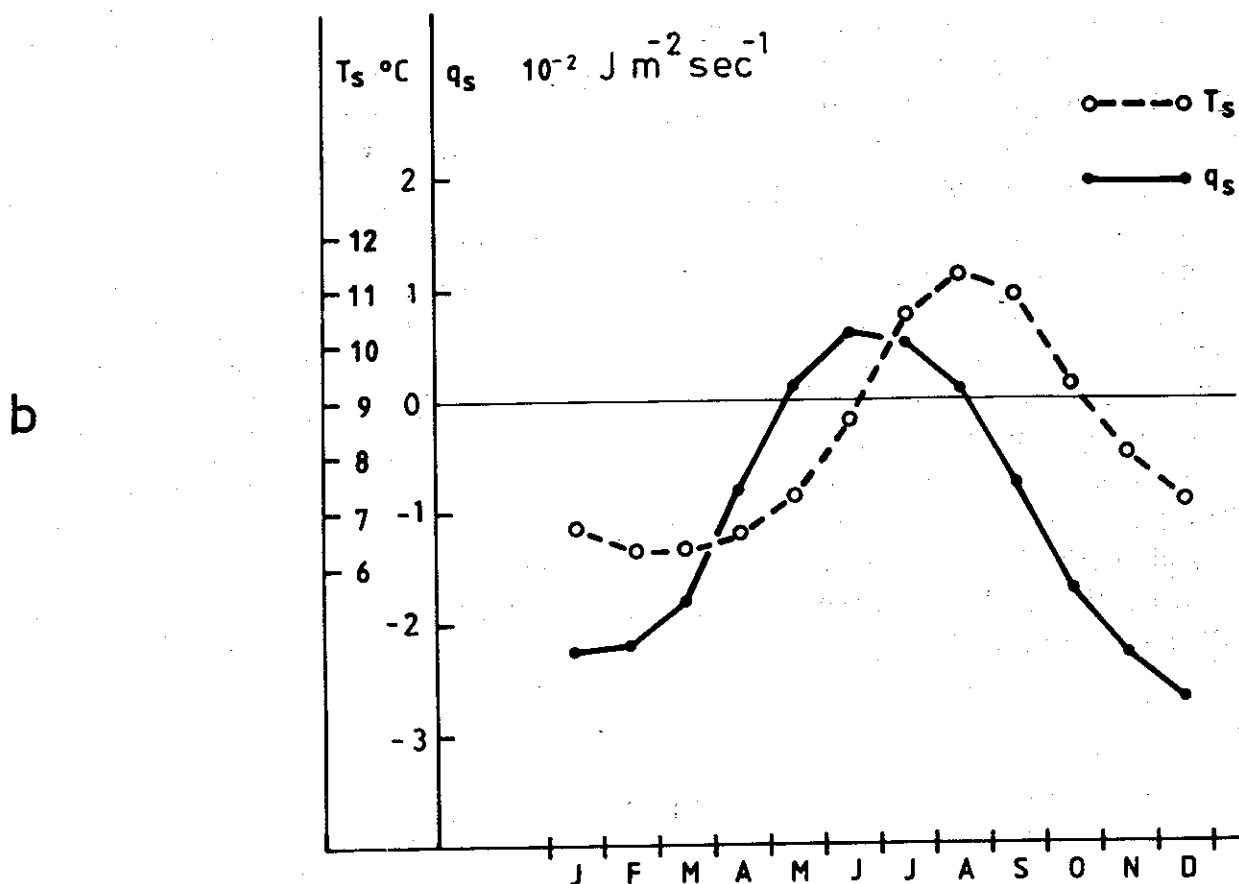
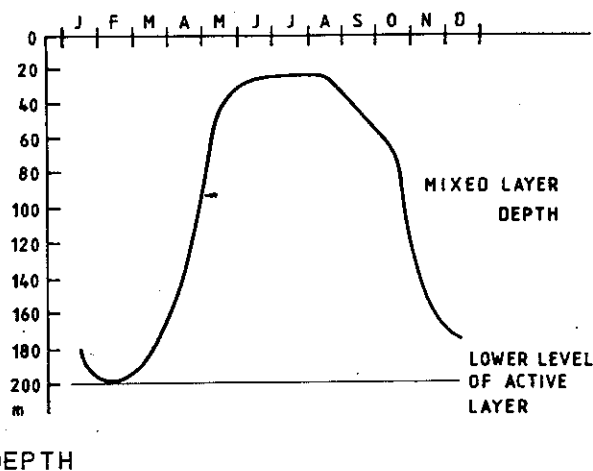


Fig. 2b. Mean seasonal variations of the heat flux q_s across the ocean atmosphere interface.

and the lower boundary of the active layer becomes so large that it starts to affect the thermal conditions of the active layer (Kitai-gorodskii 1973). The formation of the new temperature gradient determined by this temperature difference can at some depth lead to complete dampening of the vertical transport of the wind-induced turbulent energy (forced convection). The associated mixing is thereby also dampened across this gradient zone. The seasonal thermocline formed at this depth separates the upper turbulent mixed layer from the remaining part of the active layer (Fig. 3). From the beginning of the autumn the vertical mixing is determined not only by the influence of the wind but also by convection due to heat loss from the ocean surface (free convec-

tion). In this period of free and forced convective mixing the summer thermocline weakens, and around February/March the



DEPTH

Fig. 3. Mean seasonal variations of the mixed layer depth.

entire active layer becomes nearly isothermal. It is such an event of free and forced convective mixing in the mixed layer that is examined and modelled for the period 5–22 March 1967 in the next sections.

4. CASE STUDY: 5–22 MARCH

During winter time there are frequent passages of low-pressure systems in the Norwegian Sea. The monthly means of the wind speed can reach 10 m/s at this time of the year, and it is therefore of interest to see how short period variations in the upper ocean structure relate to meteorological

changes. The data base for this study consists of standard hourly marine meteorological observations, and somewhat less frequently sampled bathythermograph (BT) observations. Unfortunately, there are a few gaps where BT data are missing. The 17 day period 5–22 March 1967 was selected due to frequent passages of low-pressure systems in the region. In addition a net heat loss to the atmosphere from the sea surface occurred during this period.

The hourly observations of cloud cover, wind speed, air temperature (dry and wet) and sea surface temperature from 5–22 March 1967 are presented in Fig. 4. The mean wind speed during the 17 day period

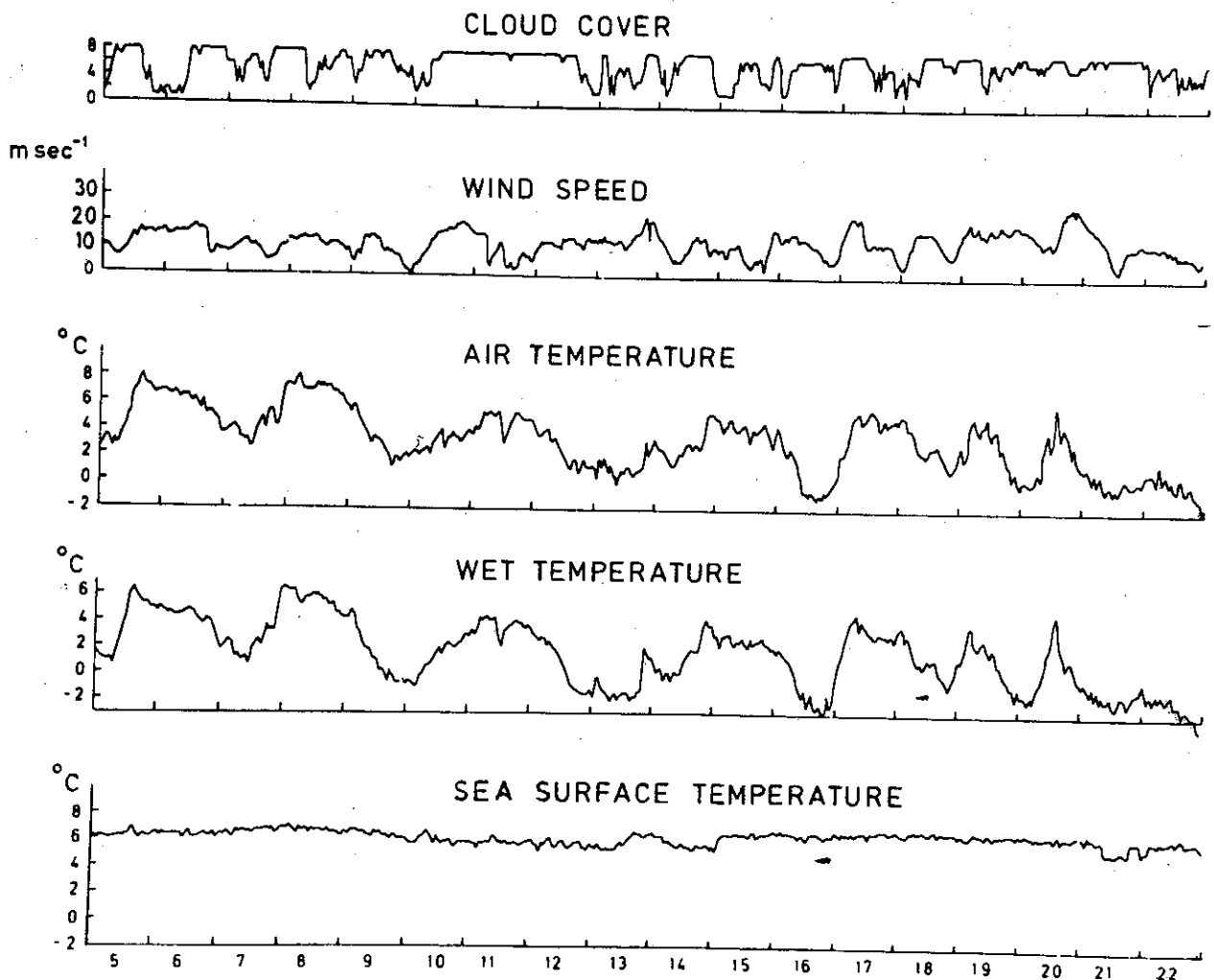


Fig. 4. Hourly meteorological observations of the cloud cover, the wind speed, the air temperature (dry and wet) and the sea surface temperature.

was about 10 m/s, but several storms with wind speeds up to 20 m/s or more were passing the area. The air temperature appears to vary between 0° C to 8° C over a period of 2 to 4 days. Similar variations are seen to take place for the wet bulb temperature. These variations are most likely associated with the passages of the cyclonic storms. The mean air temperature was about 4° C with a mean wet bulb temperature at about 1° C. During this period the sea surface temperature remained almost constant at 6° C. The atmospheric temperature fluctuations there-

fore appears to have negligible influence on the variations of the sea surface temperature. The air temperature was almost always less than that of the sea surface, causing a loss of sensible heat from the sea, as well as loss of latent heat.

The mixed layer depth (\blacktriangle) and temperature (\blacktriangle) are shown in Figs 5a and b. The mixed layer depth is here defined as the depth h to which the BT-temperature is 0.3° C less than the sea surface temperature for the same BT-cast. The temperature in the mixed layer is then set equal to the sea

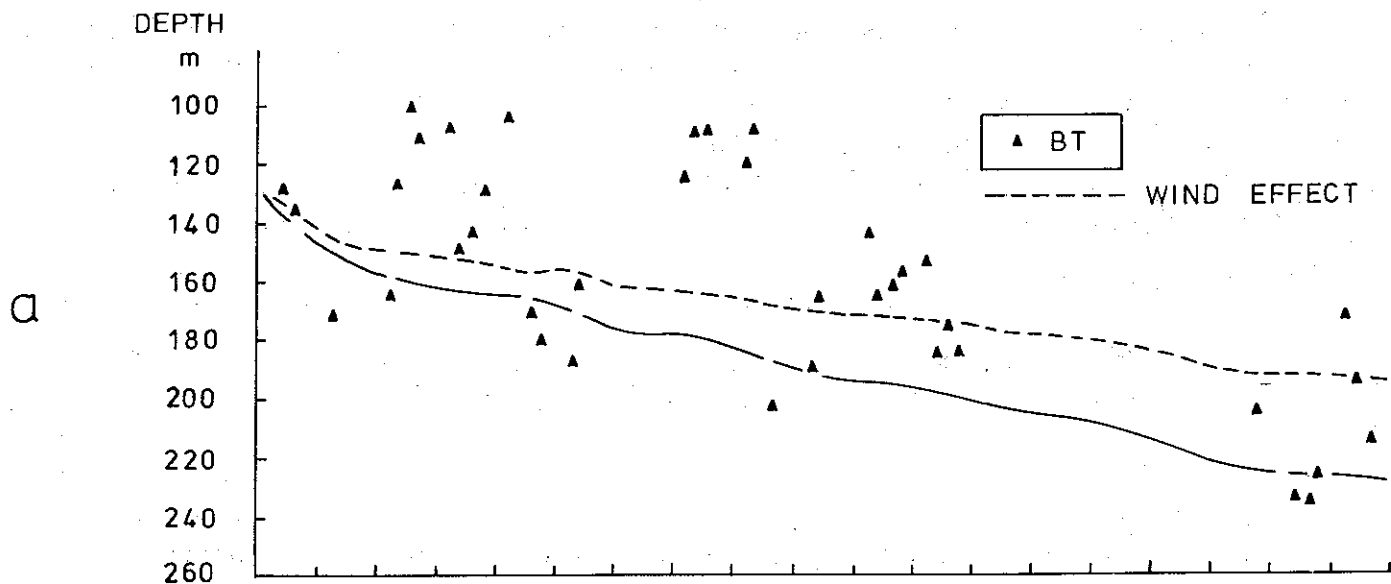


Fig. 5a. Predicted depth of the mixed layer. The observed depth (\blacktriangle) are shown for comparison.

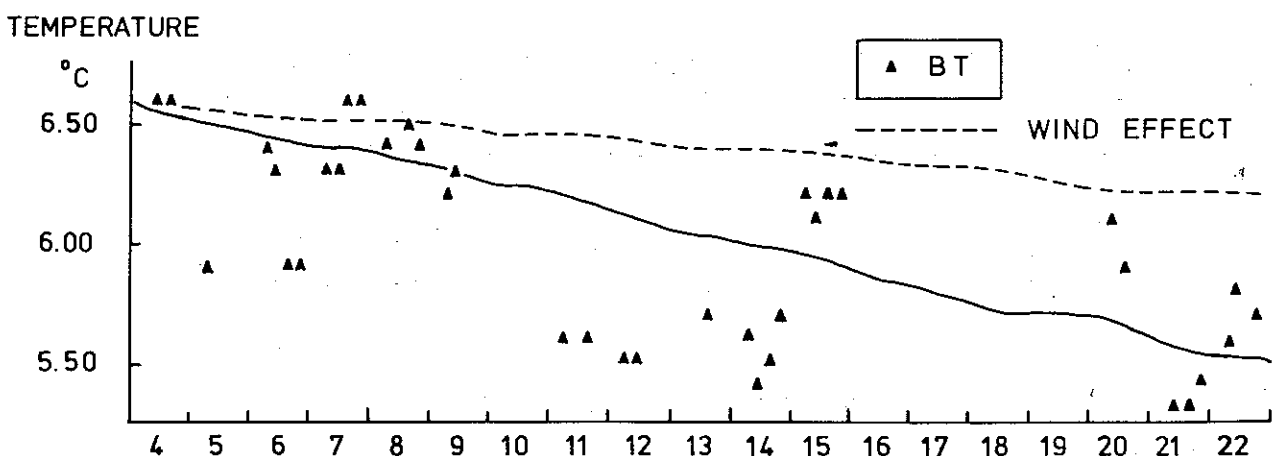


Fig. 5b. Predicted temperature of the mixed layer. The observed BT-temperature (\blacktriangle) are shown for comparison.

surface temperature T_s . The most apparent feature in Fig. 5a is the rapid fluctuations of the mixed layer—sometimes as much as 100 m in less than one day. It is difficult to recognize any dominant period in these variations due to the gaps of missing data. However, a period of 2–4 days appears to exist in correspondence with the period of the weather disturbances. The variations in the mixed layer temperature are within 5° C to 6.6° C throughout the observations (Fig. 5b). The dominant period agrees with that of the mixed layer depth but appears out of phase. A deep mixed layer is thereby much of the time warmer than a shallow mixed layer. This suggests that the changes in the structure of the mixed layer are associated with horizontal displacement of the boundary of the N.A.C. rather than an interplay of free and forced convective mixing (see Fig. 1). This is further discussed in Sect. 6. However, the speculation is partly supported by the results of the simulation of the mixed layer structure presented below.

5. THE MODEL SIMULATION

Since the mixed layer at OWS Mike is measured only by a few conventional BTs a day and much less frequent bottle casts, we are limited in testing models. There are no direct current measurements. Denman's (1973) model can be applied since it requires no current measurements and is responsive to changes on time scales of the order of one

day. It is a one-dimensional time dependent model of the upper mixed layer of the ocean driven by meteorological forcing. The model is sensitive to the rate of production of the turbulent energy by the wind stress, and to the rate of absorption with depth of the solar radiation. The available turbulent energy for mixing is assumed to be independent of depth. The temperature gradient at the base of the mixed layer has to be given. Only simple parameters available from routine meteorological measurements are required as input. For a full description of the model, see Denman (1973).

Later relevant works on one-dimensional mixed layer models, reviewed by Garwood (1979), are not taken into account, e.g. Elsberry et al. (1976) and Kim (1976), who added depth-dependent dissipation terms that helped to reduce unrealistic deepening, and De Szoeke & Rhines (1976) and Yun (1978), who added a term for the unsteadiness which became important in cases of very rapid deepening. Three-dimensional models, including horizontal and vertical advection of the mean fields such as described by Adamec et al. (1981), have also been omitted from our consideration. Such models yield more realistic simulations of the mixed layer structure in ocean areas influenced by current systems.

The deepening regime described by the dependent variables T_s , the temperature of the surface layer and h , the depth of the mixed layer, are expressed by two equations. These are

$$\frac{dT_s}{dt} = \frac{2}{\rho_0 h^2} \left\{ -\frac{(G - D)}{\alpha g} + \frac{h}{C_p} (B + H_e + H_s) \right\} + \frac{R}{C_p} (h - \gamma^{-1} \pm \gamma^{-1} e^{-\gamma h}) \quad (5.1)$$

$$\left\{ w + \frac{dh}{dt} \right\} = \frac{2 \left\{ \frac{1}{\alpha g} (G - D) + \frac{R \gamma^{-1}}{C_p} (1 - e^{-\gamma h}) \right\} - \frac{h}{C_p} \{ B + H_e + H_s + R (1 + e^{-\gamma h}) \}}{\rho_0 h (T_s - T_h)} \quad (5.2)$$

where the empirically derived variables are the surface heat fluxes of the back radiation (B), latent heat (H_e), sensible heat (H_s), the solar radiation (R) and the wind energy available for turbulent mixing (G - D). Other variables are the density ρ_0 , the acceleration of gravity g , and the specific heat capacity C_p . The extinction length γ^{-1} is taken from Jerlov (1968) to be 2.5×10^{-1} m. The value of the coefficient of thermal expansion is $\alpha = 1.67 \cdot 10^{-4} \text{ C}^{-1}$ for sea water of 10° C and 35‰. For the ratio of the potential energy increase of the water column to the downward transfer rate of turbulent energy by the wind stress, the value $m = 0.0012$, which Denman & Miyake (1973) found most suitable on the OWS Papa data, is employed. The amount of wind energy available for turbulent mixing is then expressed as

$$G - D = \rho_0 C_D \mu u^3 \quad (5.3)$$

where u is the near surface wind speed and $C_D = 1.3 \times 10^{-3}$ is the drag coefficient.

The surface heat fluxes (latent and sensible) together with the back radiation are computed using empirical formulas (Malkus 1962 and Munn 1966) where the input data are shown in Fig. 3. The fluxes are presented in Fig. 6 along with the rate of solar radiation and the wind energy available for turbulent mixing derived by Eq. (5.3).

With these data given, Eqs (5.1) and (5.2) can be solved with regard to the mixed layer depth and temperature by numerical techniques. The vertical velocity w has been set equal to zero and the time step used is 12 hours. The result is given in Figs 5a and b (solid lines). During the period 5-22 March 1967 the predicted mixed layer depth (Fig. 5a) increased smoothly from 130 m to 220 m in correspondence to strong winds and release of heat to the atmosphere from the sea surface. Moreover, the entrainment of

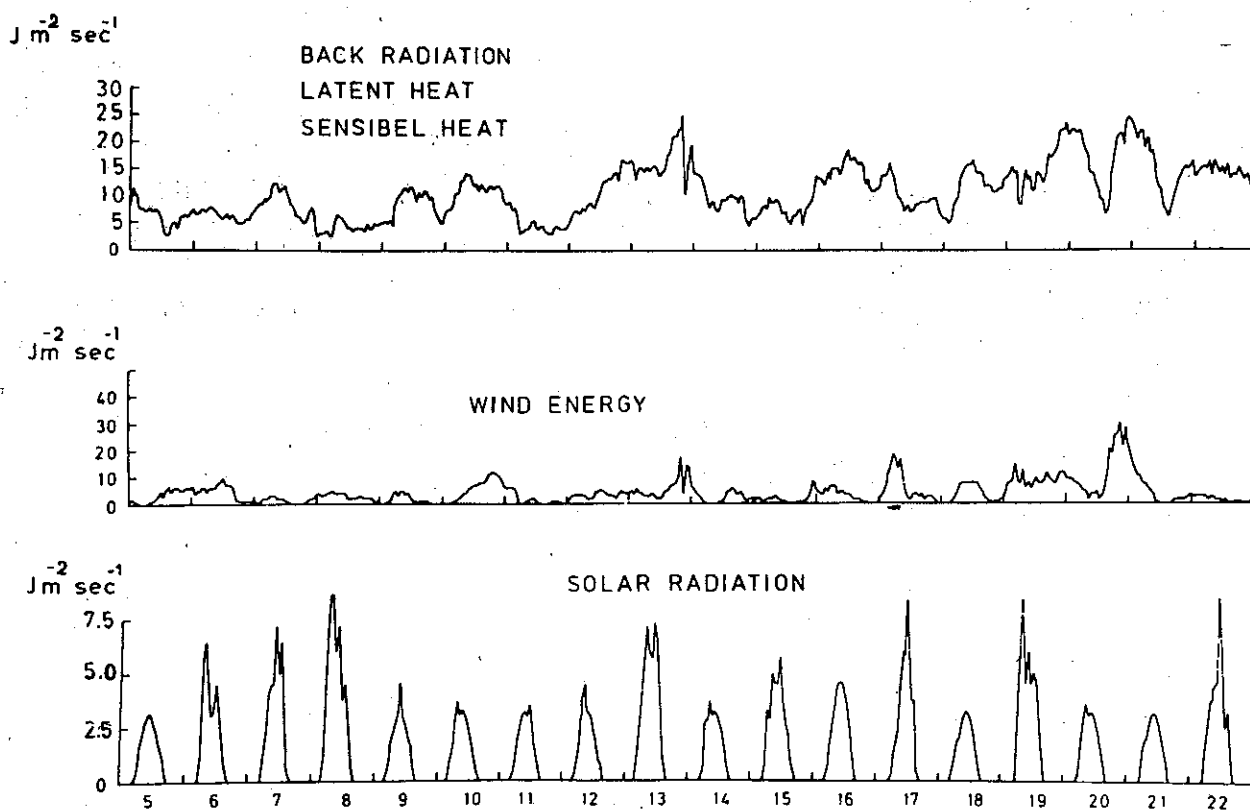


Fig. 6. Computed hourly values of the surface heat fluxes, the wind energy and the solar radiation.

cold water into the mixed layer, as well as the heat loss from the sea surface, reduced the mixed layer temperature from about 6.6°C to 5.5°C. Only a weak tendency towards more rapid increase of the mixed layer depth and decrease in the mixed layer temperature associated with the storm events can be seen.

The model can also be used to study the effect of wind mixing and free convection independently. One can assume heat balance by neglecting the radiative, latent and sensible heat transfer terms through the ocean-atmosphere interface. One can then study the growth of the mixed layer depth and decrease in the temperature due to the wind mixing. The result is seen in Figs 5a and b (dashed lines). Since the wind-induced turbulent energy available for mixing is independent of depth, as much as 60% of the increase in the mixed layer depth is accounted for by turbulent energy. Under the circumstances discussed here, with a mixed layer depth of 100–200 m, such an increase of 60% due to the wind is probably too high. The associated 20% decrease in mixed layer temperature is probably over-estimated due to the same fact. However, the choice of temperature gradient at the base of the mixed layer will also have influence on this estimate. Thus, the heat loss at the surface appears to dominate the temperature decline in the mixed layer. This is in accordance with the present knowledge of the winter cooling and convection in the Atlantic Water of the Norwegian Sea.

6. DISCUSSION AND SUMMARY

The model simulation of the mixed layer depth and temperature is in agreement with the general trend in the BTs. This suggests that the model is capable of predicting vari-

ations in the mixed layer on a weekly time scale. However, the rapid fluctuations of up to 100 m of the mixed layer depth in less than one day, every so often, are not predicted by the model. In the following, some possible causes of the discrepancies between observations and simulation are discussed.

Over the wide range of wind speeds observed during this 17 day period, it is probably not correct to keep m —the ratio of the potential energy increase to the downward transfer of turbulent energy by the wind stress—constant. For example, Turner (1969) suggests that a considerably larger value of m (0.01) would be necessary to account for the rapid storm-induced deepening of the mixed layer. Other responses in the mixed layer to passages of weather disturbances can be generation of vertical motion. With a non-zero curl in the surface wind stress field, Ekman transports are set up in the upper layer with large scale divergences and convergences. The compensating vertical motion is likely to generate fluctuations in the mixed layer temperature and depth. The combined effect of a too small m and neglect of vertical motion in the model thereby reduces the applicability of the model to simulate the true behaviour of the mixed layer during passages of storm events. However, the amplitude of these rapid fluctuations of the mixed layer are probably too large to be caused by these effects alone.

As speculated above, the most likely causes of the drastic changes in the behaviour of the mixed layer are horizontal displacements of the boundary of the N.A.C. Such displacements of the boundary of the N.A.C. may be wind-induced or due to instabilities. This conjecture is indeed supported by the infrared (IR) imagery of the southern part of the Norwegian Sea obtained on 14 May 1980 (Fig. 7). As seen in the image, the N.A.C. consists of eddies on different scales.

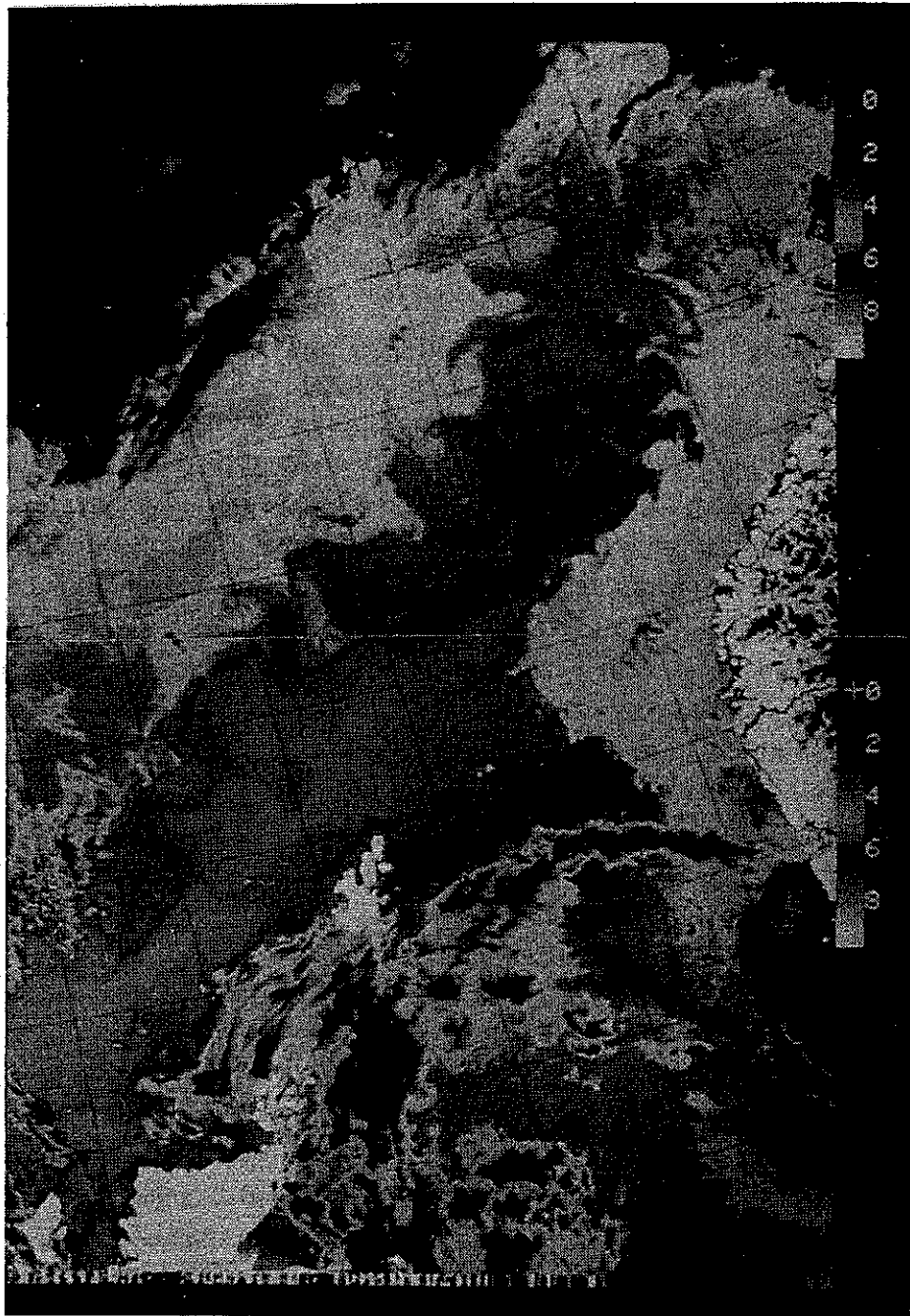


Fig. 7. IR imagery of the N.A.C. in the southern part of the Norwegian Sea obtained by TIROS-N satellite. Resolution is 1 km. (The image was produced at CMI.)

These eddies cause entrainment of colder water into the N.A.C. as well as dispersion of A.W. into the resident Norwegian Sea water and the coastal water off Norway.

The mean diameter of the eddies is in the

order of the Rossby radius of deformation, which is approximately 10 km in this region. In addition, their propagation speed is limited by the maximum phase speed of Rossby waves, expressed by McWilliams & Flierl

(1979) as

$$C = \beta L^2 \quad (6.1)$$

where β is the variation in the Coriolis parameter with latitude, and L is the Rossby radius of deformation. This speed is about 0.01 m/s at this latitude, so the total speed of the eddies will be about 0.04 m/s. The travel time for an eddy to pass OWS Mike is then approximately 2–3 days. This compares well with the major period of fluctuations seen in the BTs in Figs 5a and b. The sharp boundaries' association with these warm and cold eddies may thereby very well be responsible for the rapid changes in the structure of the mixed layer.

The discussion above therefore suggests that an increase in m , the amount of turbulent energy available for mixing, together with incorporation of vertical and horizontal advection in the model, may not be sufficient to describe the rapid fluctuations in the depth and temperature of the mixed layer. These rapid fluctuations are better explained by an alteration of water masses due to the displacement of the western boundary of the Atlantic current. Such displacement may be due to frontal waves and eddies. The propagation of these eddies appear to have time scales which correspond to the period of the rapid fluctuations.

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Probability models in precipitation climate studies

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In section 2 precipitation intensity is considered as a random variable with time as a continuous parameter. Estimates of monthly averages of precipitation probability are obtained for 280 stations.

In sections 3-7 a 3rd. order Markov model has been applied to sequences of dry and wet days. Various schemes are applied to derive probabilities for (A) the number of wet days in n successive days, (B) the duration of dry and wet runs, (C) that no more than J successive dry (wet) days occur in n days.

Sections 8-9 present a model for the distribution of precipitation sums in n days.

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INTRODUCTION

Users of information on the precipitation climate of Norway have a need for estimates of probabilities of a great variety of precipitation events, and/or related statistical quantities. In this paper the primary goal has been to show how probability theory may be used to obtain these estimates, assuming that a few basic probability models are given.

The various events which have been considered are restricted to such that characterize the precipitation climate on isolated stations, thus excluding events expressing simultaneous precipitation conditions under widely different conditions throughout the country. In this case the general structures of the models which should come into consideration are fairly well known. Problems of a more difficult nature, however, arise when the models have to be specified in detail. The systematic search for optimal models is not a subject for study in the present paper. Accordingly, the specification of

models in this work may seem, and is in fact to some extent, arbitrary. However, in spite of weaknesses in the models, they are considered good enough, for the interests of users, to compute model probabilities of some interesting events for a selection of stations. These results are tabulated in an Annex. However, with respect to more extensive results for these and other events in this article, users are asked to make specific requests.

1. THE BASIC PRECIPITATION RANDOM VARIABLE

Definitions

ω \equiv precipitation intensity as function of time: $t \rightarrow \omega(t)$; $t \in [0, \infty)$; $\omega \in [0, \infty)$

Sample space Ω :

$\Omega \equiv \{\text{all possible } \omega\text{'s}\}$

$P \equiv$ probability function on a probability field of subsets of Ω . (1)

The first random variable defined on Ω to be considered is the precipitation intensity for an arbitrarily given t , symbolized by X :

$$X = \omega(t) \tag{2}$$

This random variable has t as a parameter and takes values in $[0, \infty)$. The distributions of X ,

$$F(x) = P(X \leq x) \tag{3}$$

have shapes as illustrated by the curve marked F in Fig. 1. The jump at $x = 0$ equals the non-zero probability of dry weather. Obviously

$$\Omega = \{\omega | \omega = 0 \text{ at time } t\} \cup \{\omega | \omega > 0 \text{ at time } t\}$$

With a change in notation for the two r.h.s. disjoint subsets, this may be written more shortly

$$\Omega = \{X = 0\} \cup \{X > 0\} \tag{4}$$

Writing intersection of sets without the intersection symbol, we have

$$\begin{aligned} \{X \leq x\} &= \{X \leq x\}\Omega \\ &= \{X = 0\}\{X \leq x\} \cup \{X > 0\}\{X \leq x\} \\ &= \{X = 0\}(\{X = 0\} \cup \{0 < X \leq x\}) \cup \\ &\quad \{X > 0\}\{X \leq x\} \end{aligned}$$

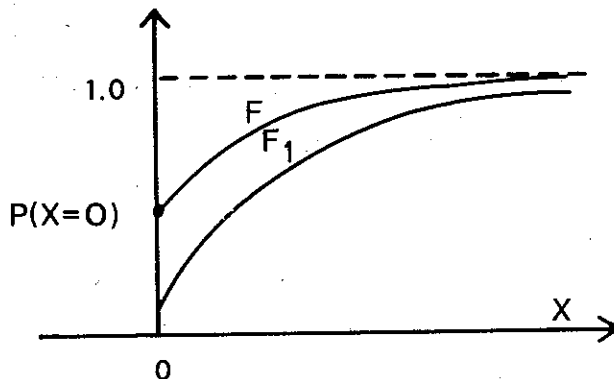


Fig. 1. A schematic illustration of probability distributions of precipitation intensity. F is the unconditional distribution and F_1 is the one relative to the subset $\{X > 0\}$.

$$= \{X = 0\} \cup \{X > 0\}\{X \leq x\}$$

The corresponding relation between probabilities becomes

$$P(X \leq x) = P(X = 0) + P(\{X > 0\}\{X \leq x\})$$

This relation may also be written

$$P(X \leq x) = P(X = 0) + P(X > 0)P(X \leq x | X > 0) \tag{5}$$

introducing the probability of $\{X \leq x\}$ relative to the subspace $\{X > 0\}$. In terms of functions of x , Eq. (5) may be written

$$F(x) = p_0 + p_1 F_1(x) \tag{6}$$

having written p_0 for $P(X = 0)$, p_1 for $P(X > 0)$, and $F_1(x)$ for $P(X \leq x | X > 0)$. Obviously

$$p_0 + p_1 = 1 \tag{7}$$

In view of this and because $F(0) = p_0$, $F(\infty) = 1$, we obtain $F_1(0) = 0$ and $F_1(\infty) = 1$ as illustrated in Fig. 1 by the curve marked F_1 .

From eq. (6) we get a corresponding relation between expectations,

$$E[X] = p_1 E[X | X > 0] \tag{8}$$

in which the r.h.s. expectation represents the expected value of precipitation intensity, not counting the cases when the intensity is zero. The fact that dry weather (zero precipitation intensity) has a non-zero probability is an essential feature of precipitation climate. As a consequence, the probability of many events, like for instance precipitation sums, will depend on the distribution of the lengths of dry and wet periods, as well as on the distribution of precipitation intensities during wet periods. Also the probability models become more complicated for the same reason.

Definition of the indicator W

$$W: X \rightarrow W(X) = \begin{cases} 0 & \text{for } X = 0 \\ 1 & \text{for } X > 0 \end{cases} \quad (9)$$

For W we have

$$E[W] = p_1 \quad (10)$$

2. TIME EVOLUTION OF X. CONTINUOUS TIME PARAMETER

Taking into account that X depends on time t as a parameter, new random variables may be defined. We shall give three of them.

Definition

$$L \equiv \int_T W dt = \text{length of precipitation}$$

time in a section T of time space

We shall restrict ourselves to a study of the expectation $E[L]$, for which we have

$$\begin{aligned} E[L] &= E\left[\int_T W dt\right] \\ &= \int_T E[W] dt = \int_T p_1 dt \end{aligned} \quad (11)$$

We have estimated $E[L]/T$ for calendar months, by the number of precipitation observations relative to the total number of observations at the main observing hours. At stations with observations at 06, 12, 18 GMT only, the estimates will be biased during summer, when most land stations have a significant daily variation in p_1 . However, this bias is small, as found already by H. Mohn (1888). Clearly, $E[L]/T$ may also be looked upon as the monthly averages of daily averages of p_1 , and therefore approximately as estimates of the daily averages of p_1 on mid-month days. The estimates are tabulated in Table 6 in the Annex.

Definition

$$D_1 \equiv \text{time until a precipitation state first comes to a stop} \quad (12)$$

$$D_0 \equiv \text{time until a non-precipitation state first comes to a stop.}$$

Adopting a stationary Markov process for the wet and dry states we have

$$P(D_1 > d_1) = e^{-d_1/E[D_1]} \quad (13)$$

$$P(D_0 > d_0) = e^{-d_0/E[D_0]}$$

It also follows that

$$\begin{aligned} P(\{W(t) = 1\}\{W(t + \Delta t) = 0\}) \\ = p_0 p_1 (1 - e^{-\Delta t/p_0 E[D_1]}) \end{aligned}$$

Putting here $\Delta t = 6$ hrs., the joint probability on the l.h.s. has been estimated from observations 6 hrs. apart. Using the tabulated numbers for p_1 in Table 6, we may solve the above equation with respect to $E[D_1]$. These numbers are tabulated in Table 7. Substituted in the first of eq. (13), we obtain the distribution for D_1 . We shall also note that when the Markov conditions above are strictly fulfilled, D_1 may also be interpreted as the lengths of precipitation "runs", i.e. the times from the last entrances into precipitating states to the first passages out of them. However, it remains to show how good an approximation the above Markov model is. This will be postponed to a later occasion, when continuous registrations of precipitation intensity will be investigated for the relatively few stations where series of such observations exist. Most likely the estimates are too high for summer months except for typical coastal stations. This is because a majority of summer showers of short duration are within the interval 12-18 GMT.

Definition

$$R \equiv \int_T X dt$$

We shall use this to get estimates of $E[X]$.

$$\begin{aligned} E[R] &= \int_T E[X] dt \\ &= \int_T p_1 E[X|X > 0] dt, \text{ using (8)} \end{aligned}$$

Under almost stationary conditions, this may be written

$$E[X|X > 0] \approx E[R]/E[L] \tag{14}$$

Table 8 contains the estimated values of $E[X|X > 0]$ having used estimated values of $E[R]$ for calendar months and the estimated values of $E[L]/T$ in Table 6 using identical observation bases.

3. PRECIPITATION VARIABLES IN DISCRETE TIME

The archives of precipitation variables in Norway are essentially of three kinds:

- (A) Continuous registrations of accumulated precipitations amounts for a few stations
- (B) Observations of precipitation conditions at fixed hours for a relatively large number of stations
- (C) Daily precipitation sums for a still higher number of stations (around 600)

Accordingly, for events connected with sequences of dry and wet days, and precipitation sums for an integer number of days, one has to use observation category (C) in order to obtain a satisfactory geographical coverage. Also it is possible that this may be a more rational way of studying events defined for periods of one day or more, rather than basing the studies on observations of time scales much less than 24 hrs.

Following up what is said above, we consider, in this and the following sections, time space as a union of successive 24 hr. intervals, which we agree to label with successive integers n . Also we now redefine the random variable X to mean:

Definition

$$X \equiv \text{precipitation sum in 24 hrs.} \tag{15}$$

This random variable depends on the parameter n . In analogy with the results of section 1, we get:

$$\begin{aligned} \Omega &= \{X = 0\} \cup \{X > 0\} \\ P(X \leq x) &= P(X = 0) + P(X > 0) \\ &\quad \times P(X \leq x | X > 0), \end{aligned}$$

or with a change in notation

$$\begin{aligned} F(x) &= P(A) + P(B)FB(x) \\ E[X] &= P(B)E[X|B] \end{aligned} \tag{16}$$

A model for $FB(x)$ will be studied in section 8.

Definition of W

$$W: X \rightarrow W(X) = \begin{cases} 0 & \text{for } X = 0 \\ 1 & \text{for } X > 0 \end{cases} \tag{17}$$

We have

$$E[W] = P(B) \tag{18}$$

4. MARKOV MODELS FOR STUDYING EVENTS DEFINED FROM SEQUENCES OF DRY AND WET DAYS

Let C denote either A or B on an arbitrary day. An arbitrary sequence of A and B 's for n successive days may then be written as the intersections

$$C_1 C_2 \dots C_n = \{C_1 C_2 \dots C_{n-1}\} C_n$$

The corresponding relations between prob-

abilities then becomes

$$P(C_1 C_2 \dots C_n) = P(\{C_1 C_2 \dots C_{n-1}\} C_n) \\ = P(C_1 C_2 \dots C_{n-1}) \\ \times P(C_n | C_1 C_2 \dots C_{n-1})$$

Here $P(C_n | C_1 C_2 \dots C_{n-1})$ formally expresses the probabilities of $C = A$ or $C = B$ on the n -th day as depending on a knowledge of the distribution of dry and wet days on the previous $n - 1$ days. When no such dependence exists, we have

$$P(C_n | C_1 C_2 \dots C_{n-1}) = P(C_n) \text{ for any } n$$

When there is a dependence on history, we shall assume that it is of the following kind:

$$P(C_n | C_1 C_2 \dots C_{n-1}) \\ = P(C_n | C_{n-k} C_{n-k+1} \dots C_{n-1}) \quad (19)$$

for fixed k and any value of n (Markov model of order k). We note that

$$P(A_n | C_{n-k} \dots C_{n-1}) \\ + P(B_n | C_{n-k} \dots C_{n-1}) = 1$$

Therefore, there are altogether 2^k independent transition probabilities on any day n . In Table 1 we give estimates of $P(A_n)$ and the particular transition probabilities $P(A_n | A_{n-k} \dots A_{n-1})$ for $k = 1, 2, 3$ at a station in SE-Norway. The estimates are based upon 82 years of observations.

In this paper orders higher than 3 will not be considered, Eidsvik (1979), Bjørgum (1945), Bjørgum & Fjordholm (1946). As a

Table 1

Transition probabilities $P(A_n | A_{n-k} \dots A_{n-1})$ for sequences of dry and wet days as function of order k of Markov chain at a station in south-eastern Norway.

Month	Indep.	$k = 1$	$k = 2$	$k = 3$
March	.7400	.8334	.8453	.8588
August	.5651	.7114	.7411	.7617

general rule, the shorter the periods from which the events are defined, the better will the Markov model work.

5. THE NUMBER L OF PRECIPITATION DAYS IN N SUCCESSIVE DAYS

This random variable may be written $L = \sum_{i=1}^N W_i$. Hence

$$E[L] = \sum_{i=1}^N E[W_i] = \sum_{i=1}^N P(B_i) \quad (20)$$

Under almost stationary conditions, this may be written

$$E[L] \approx NP(B) \quad (21)$$

In order to derive the distribution $f(l, N) \equiv P(L = l)$, $l = 0, 1, \dots, N$, we shall adopt a 3rd order Markov model for sequences of precipitation and non-precipitation days. We have

$$\Omega = A_{i-2} A_{i-1} A_i \cup A_{i-2} A_{i-1} B_i \\ \cup A_{i-2} B_{i-1} A_i \cup A_{i-2} B_{i-1} B_i \\ \cup B_{i-2} A_{i-1} A_i \cup B_{i-2} A_{i-1} B_i \\ \cup B_{i-2} B_{i-1} A_i \cup B_{i-2} B_{i-1} B_i \quad (22)$$

and

$$\Omega = A_i \cup B_i \quad (23)$$

the latter having been used already in section 3.

A consequence of (19) is

$$P(Z_{n-1} C_n) \\ = P(Z_{n-1}) P(C_n | C_{n-3} C_{n-2} C_{n-1}) \quad (24)$$

if

$$Z_{n-1} = (\cup_J (C_1 C_2 \dots C_{n-4}) J) \\ \times C_{n-3} C_{n-2} C_{n-1} \quad (25)$$

in other words, if Z_{n-1} is defined on the first $n - 1$ outcomes, but such that the three last

intervals have a given sequence, $C_{n-3}C_{n-2}C_{n-1}$, of outcomes.

Let Z represent an arbitrary non-empty subset of Ω . Then

$$\begin{aligned}
 Z A_{n-1} A_n A_{n+1} &= Z A_{n-2} A_{n-1} A_n A_{n+1} \\
 &\cup Z B_{n-2} A_{n-1} A_n A_{n+1} \\
 Z A_{n-1} A_n B_{n+1} &= Z A_{n-2} A_{n-1} A_n B_{n+1} \\
 &\cup Z B_{n-2} A_{n-1} A_n B_{n+1} \\
 Z A_{n-1} B_n A_{n+1} &= Z A_{n-2} A_{n-1} B_n A_{n+1} \\
 &\cup Z B_{n-2} A_{n-1} B_n A_{n+1} \\
 Z A_{n-1} B_n B_{n+1} &= Z A_{n-2} A_{n-1} B_n B_{n+1} \\
 &\cup Z B_{n-2} A_{n-1} B_n B_{n+1} \\
 Z B_{n-1} A_n A_{n+1} &= Z A_{n-2} B_{n-1} A_n A_{n+1} \\
 &\cup Z B_{n-2} B_{n-1} A_n A_{n+1} \\
 Z B_{n-1} A_n B_{n+1} &= Z A_{n-2} B_{n-1} A_n B_{n+1} \\
 &\cup Z B_{n-2} B_{n-1} A_n B_{n+1} \\
 Z B_{n-1} B_n A_{n+1} &= Z A_{n-2} B_{n-1} B_n A_{n+1} \\
 &\cup Z B_{n-2} B_{n-1} B_n A_{n+1} \\
 Z B_{n-1} B_n B_{n+1} &= Z A_{n-2} B_{n-1} B_n B_{n+1} \\
 &\cup Z B_{n-2} B_{n-1} B_n B_{n+1}
 \end{aligned}
 \tag{26}$$

where, in consequence of (22) the union of the eight left members equals Z .

That the sets above are identical is easily recognized by using $A_{n-2} \cup B_{n-2} = \Omega$. The probabilities of the events on the left sides of eq. (26) are obtained by adding the probabilities of the events on the right sides. Let now as a special case $Z = Z_{l,n+1} \equiv$ union of all sample points with exactly l precipitation days in the $n + 1$ first intervals. Then in the r.h.s. members of eq. (26), $Z_{l,n+1}$ may be replaced by $Z_{l,n}$ in the relations ending with a non-precipitation day, and with $Z_{l-1,n}$ in the relations where an additional precipita-

tion day in the interval $n + 1$ adds to the previous number $l - 1$ in the first n intervals. We define $f_k(l, n)$, $p_k A(n)$, and $p_k B(n)$ by

$$\begin{aligned}
 f_1(l, n) &\equiv P(Z(l, n) A_{n-2} A_{n-1} A_n) \\
 f_2(l, n) &\equiv P(Z(l, n) A_{n-2} A_{n-1} B_n) \\
 p_1 A(n) &\equiv P(A_{n+1} | A_{n-2} A_{n-1} A_n) \\
 p_1 B(n) &= 1 - p_1 A(n) \\
 p_2 A(n) &\equiv P(A_{n+1} | A_{n-2} A_{n-1} B_n) \\
 p_2 B(n) &= 1 - p_2 A(n)
 \end{aligned}$$

Then, by applying eq. (24) to the right members of eq. (26), we get:

$$\begin{aligned}
 f_1(l, n + 1) &= p_1 A(n) f_1(l, n) \\
 &\quad + p_5 A(n) f_5(l, n) \\
 f_2(l, n + 1) &= p_1 B(n) f_1(l - 1, n) \\
 &\quad + p_5 B(n) f_5(l - 1, n) \\
 &\quad \vdots \\
 f_8(l, n + 1) &= p_4 B(n) f_4(l - 1, n) \\
 &\quad + p_8 B(n) f_8(l - 1, n) \\
 f(l, n) &= \sum_{k=1}^8 f_k(l, n); l = 0, 1, \dots, n.
 \end{aligned}
 \tag{27}$$

Equations (27) represent an iteration scheme from which $f(l, n)$ may be obtained for any n , given the initial fields $f_k(0, 0)$ and the transition probabilities $p_k A(i)$; $i = 1, 2, \dots, N$. We have

$$\begin{aligned}
 f_1(0, 0) &= P(A_{-2} A_{-1} A_0) \\
 f_2(0, 0) &= P(A_{-2} A_{-1} B_0) \\
 &\quad \vdots \\
 f_8(0, 0) &= P(B_{-2} B_{-1} B_0)
 \end{aligned}
 \tag{28}$$

Table 9 contains for $N = 30$ computed values of $F(l, n) \equiv \sum_{j=0}^l f(j, N)$ for selected values of

1 for 33 stations. More extensive tables may be obtained on request.

6. DURATION OF PRECIPITATION RUNS

Given that a precipitation run has started, the probability that it will last at least d days is given by

$$G(d) \equiv P(DI \geq d) = P(A_0 B_1 B_2 \dots B_d) / P(A_0 B_1) \quad (29)$$

In our model, this equation becomes:

$$G(d) = \{P(A_0) \cdot P(B_1|A_0) \cdot P(B_2|A_0 B_1) \cdot P(B_3|A_0 B_1 B_2) \cdot P(B_4|B_1 B_2 B_3) \dots P(B_d|B_{d-3} B_{d-2} B_{d-1})\} / P(A_0 B_1)$$

giving

$$\begin{aligned} G(1) &= 1 \\ G(2) &= P(B_2|A_0 B_1) \\ G(3) &= G(2)P(B_3|A_0 B_1 B_2) \end{aligned} \quad (30)$$

$$G(d) = G(3) \prod_{n=4}^d P(B_n|B_{n-3} B_{n-2} B_{n-1}); d > 3$$

Under almost stationary conditions this may be written more simply as

$$\begin{aligned} G(1) &= 1 \\ G(2) &= P(B|AB) \\ G(d) &= G(2)P(B|ABB) \\ &\quad \times [P(B|BBB)]^{d-3}; d \geq 3 \end{aligned} \quad (31)$$

For non-precipitation run $D0$, the corresponding equations are, in the stationary case,

$$\begin{aligned} F(1) &= 1 \\ F(2) &= P(A|BA) \\ F(d) &= F(2)P(A|BAA) \\ &\quad \times [P(A|AAA)]^{d-3}; d \geq 3 \end{aligned} \quad (32)$$

where $F(d) \equiv P(D0 \geq d)$.

We note that

$$P(d_1 \leq D1 < d_2) = G(d_1) - G(d_2) \quad (33)$$

$$P(d_1 \leq D0 < d_2) = F(d_1) - F(d_2) \quad (34)$$

In order to exemplify the use of the above formula, for two stations, whose locations may be found from List B, we have computed $100 \cdot F(D0 \equiv d)$ for $d = 1$ up to 15 days, as tabulated in Table 2.

7. THE EVENT Z_n , THAT THE NUMBER OF SUCCESSIVE DRY DAYS IN n SUCCESSIVE DAYS DOES NOT EXCEED J

This is an event which may cause considerable harm to activities for which a minimum number of successive dry days within periods of given lengths is essential. For large values of J , the complementary events \bar{Z}_n defined by $Z_n \cup \bar{Z}_n = \Omega$, represent outcomes of serious drought within periods of given lengths.

In the following we shall develop an iterative scheme to derive $P(Z_n; J)$ for increasing values of n . The probability of the complementary events, that at least $J + 1$ successive days occur in a period of n days, is next found from $P(\bar{Z}_n; J) = 1 - P(Z_n; J)$. We note that $P(Z_n) = 1$ when $n \leq J$, whereby n may be taken as $n \geq J + 1$.

Let i represent the number of days we have to go back from day n in order to find the first wet day. With $\Omega^{(i)}$ symbolizing the corresponding events we find

$$\begin{aligned} \Omega^{(0)} &= B_n \\ \Omega^{(1)} &= B_{n-1} A_n \\ \Omega^{(2)} &= B_{n-2} A_{n-1} A_n \\ \Omega^{(J)} &= B_{n-J} A_{n-J+1} \dots A_n \end{aligned} \quad (35)$$

Obviously

$$\Omega = \bigcup_{i=0}^J \Omega^{(i)} \cup \left(\bigcup_{i=J+1}^{\infty} \Omega^{(i)} \right)$$

Table 2. Model probability estimates that the total lengths of unbroken periods of dry days are equal to or greater than 2, 3, . . . , 15 days for stations 1252 and 5056 in List B in the Annex, respectively. Unit: 1/1000.

	2	3	4	5	6	7	8	9	10	11	12	13	14	15
J	68	47	37	29	23	18	14	11	9	7	5	4	3	3
F	72	52	42	35	29	24	19	16	13	11	9	7	6	5
M	75	59	51	43	37	32	28	24	20	17	15	13	11	9
A	76	61	51	43	36	30	25	21	18	15	13	11	9	7
M	69	51	43	36	30	25	20	17	14	12	10	8	7	6
J	66	48	38	31	25	20	17	13	11	9	7	6	5	4
J	63	44	33	25	19	14	11	8	6	5	3	3	2	2
A	61	42	32	25	19	14	11	8	6	5	4	3	2	2
S	63	47	37	29	22	17	14	11	8	6	5	4	3	2
O	66	46	37	29	23	18	14	11	9	7	6	4	3	3
N	66	45	36	29	23	18	15	12	9	7	6	5	4	3
D	64	44	35	27	21	16	13	10	8	6	5	4	3	2
J	57	39	29	22	16	12	9	7	5	4	3	2	2	1
F	60	42	31	23	17	13	9	7	5	4	3	2	2	1
M	60	43	33	25	19	15	11	8	6	5	4	3	2	2
A	60	40	28	20	14	10	7	5	4	3	2	1	1	1
M	66	46	35	27	21	16	12	9	7	6	4	3	2	2
J	56	39	30	23	18	14	11	8	7	5	4	3	2	2
J	55	38	27	19	14	10	7	5	4	3	2	1	1	1
A	57	37	26	19	14	10	7	5	4	3	2	1	1	1
S	56	32	22	15	10	7	5	3	2	2	1	1	0	0
O	55	36	24	16	11	7	5	3	2	1	1	1	0	0
N	54	35	24	16	11	7	5	3	2	2	1	1	0	0
D	57	38	27	19	13	9	7	5	3	2	2	1	1	1

Using $Z_n = Z_n \Omega$ and denoting $Z_n \Omega^{(i)}$ by $Z_n^{(i)}$, we obtain

$$Z_n = \bigcup_{i=0}^J Z_n^{(i)} \cup \text{empty set}, \tag{36}$$

since Z_n according to its definition is disjoint to $\bigcup_{j=1}^{\infty} \Omega^{(j)}$. Furthermore we see that

$$\begin{aligned} Z_{n+1}^{(0)} &= Z_n^{(0)} B_{n+1} \cup \dots \cup Z_n^{(J)} B_{n+1} \\ Z_{n+1}^{(1)} &= Z_n^{(0)} A_{n+1} \\ &\vdots \\ Z_{n+1}^{(J)} &= Z_n^{(J-1)} A_{n+1} \end{aligned} \tag{37}$$

The probability relations corresponding to (37) and (36) become in a 3rd order Markov model:

$$\begin{aligned} P_{n+1}^{(0)} &= P_n^{(0)} \cdot P(B_{n+1} | B_n) \\ &+ P_n^{(1)} \cdot P(B_{n+1} | B_{n-1} A_n) \\ &+ P_n^{(2)} \cdot P(B_{n+1} | B_{n-2} A_{n-1} A_n) \\ &+ \sum_{i=3}^J P_n^{(i)} \cdot P(B_{n+1} | A_{n-2} A_{n-1} A_n) \\ P_{n+1}^{(1)} &= P_n^{(0)} \cdot P(A_{n+1} | B_n) \\ P_{n+1}^{(2)} &= P_n^{(1)} \cdot P(A_{n+1} | B_{n-1} A_n) \\ P_{n+1}^{(3)} &= P_n^{(2)} \cdot P(A_{n+1} | B_{n-2} A_{n-1} A_n) \\ P_{n+1}^{(i)} &= P_n^{(i-1)} \cdot P(A_{n+1} | A_{n-2} A_{n-1} A_n); \\ &\quad i = 4, 5, \dots, J \\ P_n &= \sum_{i=0}^J P_n^{(i)} \end{aligned} \tag{38}$$

Eq. (38) represents an iteration scheme from which the values of $P_n^{(i)}$ for any n may be obtained, knowing their initial values, and given estimates of the transition probabilities appearing in the equations. Since Z_n is cer-

tain when $n \leq J$, the initial n may be taken to be $J + 1$, for which we have the initial set of probabilities:

$$\begin{aligned} P_{J+1}^{(0)} &= P(B_{J+1}) \\ P_{J+1}^{(1)} &= P_{J+1}^{(0)} \cdot P(A_{J+1} | B_J) \\ P_{J+1}^{(2)} &= P_{J+1}^{(1)} \cdot P(A_{J+1} | B_J) \\ P_{J+1}^{(3)} &= P_{J+1}^{(2)} \cdot P(A_{J+1} | B_{J-1} A_J) \\ P_{J+1}^{(i)} &= P_{J+1}^{(i-1)} \cdot P(A_{J+1} | A_{J-2} A_{J-1} A_J); \\ &\quad i = 4, 5, \dots, J \end{aligned} \tag{40}$$

Statistics of the above kind may be obtained on request for arbitrary values of J and n (within certain limits), for a number of stations. Below we include, by way of example, a few results in Tables 3 and 4.

8. THE DISTRIBUTION OF PRECIPITATION SUMS, R

$$R \equiv \sum_{i=1}^n X_i \tag{41}$$

Decomposing Ω into $\Omega = \{L = 0\} \cup \{L = 1\} \cup \dots \cup \{L = n\}$, we obtain

$$\begin{aligned} \{R \leq r\} &= \Omega \{R \leq r\} = \{L = 0\} \cup \{L = 1\} \\ &\quad \times \{R \leq r\} \cup \dots \cup \{L = n\} \{R \leq r\} \end{aligned}$$

Taking probabilities of these events, gives

$$\begin{aligned} P(R \leq r) &= P(L = 0) + \sum_{i=1}^n P(L = i) \\ &\quad \times P(R \leq r | L = i), \end{aligned}$$

or in terms of functions of l and r

$$F(r; n) = f(0, n) + \sum_{l=1}^n f(l, n) F_l(r; n) \tag{42}$$

Since $F(\infty; n) = 1$, $F(0; n) = f(0, n)$ and

Table 3

Model probability estimates that no more than J successive dry days occur in n days on Station 2780, List B. October. Unit: 1/1000.

n		J = 1								
2-10		455	353	277	217	170	133	104	82	64
11-20	50	39	31	24	19	15	12	9	7	6
n		J = 2								
3-10			582	506	430	368	315	270	231	198
11-20	169	145	124	106	91	78	67	57	49	42
21-30	36	31	26	22	19	16	14	12	10	9
n		J = 3								
4-10				676	616	557	497	447	402	361
11-20	324	292	262	235	211	190	171	153	138	124
21-30	111	100	90	81	72	65	59	53	47	42
31-40	38	34	31	28	25	22	20	18	16	15

$\sum_{l=0}^n f(l, n) = 1$ we obtain $F_1(0; n) = 0$ and $F_1(\infty; n) = 1$. The point probabilities $f(l, n)$, as demonstrated in section 5, may be found from the 3rd-order Markov model described earlier. With regard to the distributions F_1 we now adopt the model

$$F_1(r; n) = \text{Gam}(r; \alpha_{l,n}, \beta_{l,n})$$

$$= \int_0^{r/\beta_{l,n}} \frac{U^{\alpha_{l,n}}}{\alpha_{l,n}!} e^{-U} dU;$$

$$l = 1, 2, \dots, n \quad (43)$$

as an approximation. This hypothesis is supported by observations for those values of l and n where observation series are long

Table 4

Probability that at least J + 1 successive dry days occur in n days. Station 2780, List B. Unit: 1/1000.

March-April		June-July	
n	J = 29	n	J = 22
30	6,1	23	2,5
59	32,9	55	14,1

enough to permit a significant testing of the model. For large values of n it would, however, be impossible to estimate all 2n parameters from available observation series. In this case it is therefore essential that the number of parameters, by further development of the model, should be reduced drastically. In what follows we shall describe how this may be accomplished.

From the model representation of $F_1(r; n)$ we obtain

$$\beta_{l,n}(\alpha_{l,n} + 1) = E[R|L = l] \quad (44)$$

$$\beta_{l,n}^2(\alpha_{l,n} + 1) = \text{VAR}[R|L = l] \quad (45)$$

Now, R is the sum $X^{(1)} + X^{(2)} + \dots + X^{(l)}$ of l daily precipitation sums without any ordering relative to the n different days. We may therefore write

$$E[R|L = l] = lE[X|B, L = l] \quad (46)$$

$$\text{VAR}[R|L = l] = \text{VAR}[X|B, L = l] \cdot (1 + Q) \quad (47)$$

where

$$Q = 2 \sum_{d=1}^{n-1} P(W_i = 1, W_{i+d} = 1 | L = 1) \cdot (n - d) \cdot \text{cor}[X_i X_{i+d} | W_i = 1, W_{i+d} = 1, L = 1] \quad (48)$$

At first sight the estimation problem has now become much worse, when using the r.h.s. expressions in (46), (47). For, as we shall soon see, $E[X|B, L = 1]$ and $\text{VAR}[X|B, L = 1]$ have a real dependency on l . This alone gives $2n$ quantities to be estimated as before. Furthermore, whereas the conditional joint probability function, at least in theory, is determined from our model for sequences of dry and wet days, we have in addition got conditional correlations to be estimated for $d = 1, 2, \dots, n - 1$ for each $l = 2, \dots, n$, i.e. altogether $2(n - 1)^2$ correlations. In the subparagraphs (A)–(C) below, we shall show how a real reduction in the numbers of quantities to be estimated may be accomplished, without having to make the model too unrealistic.

(A) Stratification of wet days

We introduce 3 types of wet days, B_1, B_2, B_3 , as shown below:

$$AB_1A, AB_2B \text{ or } BB_2A, \text{ and } BB_3B \quad (49)$$

We shall then find that the means and standard deviations of daily precipitation amounts differ systematically for these 3 types. For station 2370 we find, for instance, the following estimates in the month of October

$$\begin{aligned} E[X|B_1] &= 5.87 \text{ mm;} \\ E[X|B_2] &= 8.00 \text{ mm;} \\ E[X|B_3] &= 11.94 \text{ mm} \\ SD[X|B_1] &= 6.97 \text{ mm;} \end{aligned} \quad (50)$$

$$SD[X|B_2] = 9.47 \text{ mm;} \quad (51)$$

$$SD[X|B_3] = 12.00 \text{ mm}$$

The same ordering with respect to the magnitudes of these quantities has been found in all months at all 33 stations except one, station 165. Assuming that the first and second moments of X , given the B_k , are not influenced by a further knowledge of the numbers L_k of B_k we obtain

$$E[X|B, L = 1] = \sum_{k=1}^3 M_k E[X|B_k] \quad (51)$$

$$E[X^2|B, L = 1] = \sum_{k=1}^3 M_k E[X^2|B_k] \quad (52)$$

where the weight M_k is the expected value of the number L_k of wet days of type k relative to the total number l of wet days in the subset,

$$\begin{aligned} M_k &= E[L_k | L = 1] / l \\ &= \sum_{l_1+l_2+l_3=1} l_k \\ &\quad \times P(L_1 = l_1, L_2 = l_2, L_3 = l_3) / l \end{aligned} \quad (53)$$

The point probabilities of exactly l_1 wet days of the first kind, l_2 of the second, and l_3 of the third in n days may be found using an iteration scheme of a type completely analogous to the one used in Section 5. The corresponding expression for $\text{VAR}[X|B, L = 1]$ is found from

$$\text{VAR}[X|B, L = 1] = E[X^2|B, L = 1] - E^2[X|B, L = 1]$$

after substituting from (51), (52):

$$\begin{aligned} \text{VAR}[X|B, L = 1] &= \sum_{k=1}^3 M_k E[X^2|B_k] \\ &\quad - \left(\sum_{k=1}^3 M_k E[X|B_k] \right)^2 \end{aligned} \quad (54)$$

Using the estimates in (50), and the com-

puted values of M_k for $l = 1, 2, \dots, 30$, with $n = 30$ found from a 3rd order Markov model, we obtain the results in Table 5.

Obviously with increasing l , the chances increase that there will be more wet days of type 3 relative to the number of wet days of type 2 and 1, and more wet days of type 2 relative to the number of wet days of type 1. This, taken together with the ordering relations in (50), in short explains the reasons behind the result why $E[X|B, L = l]$ and $SD[X|B, L = l]$ increase for increasing values of l .

(B) *Simplification of $P(W_i = 1, W_{i+d} = 1|L = l)$ in Eq. (48)*

The expression for Q in eq. (48) implies that stationarity is assumed within our subsets. The subscripts i and $i + d$ indicate two different days picked out at random. We now assume, as an approximation, that the joint probability in question is independent of d , in which case

$$P(W_i = 1, W_{i+d} = 1|L = l) = \frac{1}{n} \cdot \frac{l-1}{n-1} \quad (55)$$

(C) *Simplification of the correlation function*

The correlation function appearing in (48) is next assumed to be independent of l :

$$\begin{aligned} \text{cor}[X_i X_{i+d} | W_i = 1, W_{i+d} = 1, L = l] \\ = \text{cor}[X_i X_{i+d} | W_i = 1, W_{i+d} = 1] \end{aligned} \quad (56)$$

Although we do not see any obvious reason why this function should depend strongly on l , we do best in not considering (56) as an exact relation. What is important in our connection is the fact that (56) taken together with (55) simplifies the estimation problem drastically. By substituting from (55), (56) into (48), we get

$$\begin{aligned} Q = l(l-1) \cdot \sum_{d=1}^{n-1} (n-d) \\ \times \text{cor}[X_i X_{i+d} | W_i = 1, W_{i+d} = 1] \\ \cdot \frac{2}{n(n-1)} \end{aligned} \quad (57)$$

where the sum is now just a parameter, C , independent of d and l :

$$\begin{aligned} C \equiv \sum_{d=1}^{n-1} (n-d) \text{cor}[X_i X_{i+d} | W_i = 1, W_{i+d} = 1] \\ \cdot \frac{2}{n(n-1)} \end{aligned} \quad (58)$$

Then (47) assumes the form:

$$\begin{aligned} \text{VAR}[R|L = l] = \text{VAR}[X|B, L = l] \\ \times (1 + l(l-1)C) \end{aligned} \quad (59)$$

Table 5

Model estimates of expectations and standard deviations of daily precipitation amounts on days of precipitation in subsets of l wet days out of 30. Station 2370, List B. October. Unit: 1/1000.

l	$E[X B, L = l]$									
1-10	61	68	72	75	78	80	82	84	86	88
11-20	89	91	92	93	95	96	97	99	100	101
21-30	103	104	105	107	109	110	112	114	116	119
l	$SD[X B, L = l]$									
1-10	73	82	87	90	93	95	97	99	100	102
11-20	103	104	105	106	107	108	109	110	111	111
21-30	112	113	114	115	116	117	117	119	119	121

For the unconditional variance of the precipitation sums of n days we have

$$\begin{aligned} \text{VAR}[R; n] &= \sum_{l=1}^n f_l E[R^2|L=l] \\ &\quad - \left(\sum_{l=1}^n f_l E[R|L=l] \right)^2 \\ &= \sum_{l=1}^n f_l (\text{VAR}[R|L=l] \\ &\quad + E^2[R|L=l]) - \left(\sum_{l=1}^n f_l E[R|L=l] \right)^2 \quad (60) \end{aligned}$$

Therefore this variance depends linearly on C . Instead of estimating C directly from the expression in (58), which would be very difficult, we may therefore estimate C by taking a value for it that will make the value of $\text{VAR}[R; n]$ as obtained from (60) equal to its directly estimated value from observations.

In conclusion, in addition to the parameters in the Markov model for sequences of dry and wet days, we have to estimate the six parameters in (50) and the variance $\text{VAR}[R; n]$ to get estimates of $E[R|L=l]$ and $\text{VAR}[R|L=l]$ from which the parameters of the subset Gammadistributions are obtained, using (44), (45).

9. REMARKS ON THE USE OF THE MODEL IN SECTION 8

In estimating the parameters for the model of the previous section, observations have been grouped in twelve groups, one for each calendar month, and no attempt has been made to distinguish between observations belonging to the same group.

This implies that the climatological trends within calendar months have been con-

sidered as sufficiently small to justify considering conditions as quasi-stationary within periods of the lengths of months. However, in order to eliminate as much as possible influences from the trends, which, although small, do exist, the results obtained for the probability distributions of precipitation sums should be interpreted as valid for periods centered around mid-month days. Results for other periods may then be obtained by interpolation.

The distribution function $F(R; n)$ may then be computed for calendar months for values of $n \leq$ length of calendar month, and for selected values of R , as desired. In this article we shall just reproduce F for $n = 30$ and for values of R given by $100 \cdot R/E'[R] = 2, 5, 10, 15, 25, 30, 50, 75, 100, 150, 200, 250, 300,$ and 400 , where $E'[l]$ is an estimate of $E[R; 30]$. The numbers are tabulated in Table 10 in the Annex.

In Fig. 2, points $(F, R/E'[R])$ for the above values of R as found for station 3845 have been plotted and connected with unbroken lines. In the same figure crossmarks give a selection of corresponding points as obtained by frequency countings from 82 years of observations.

The degree of correspondence between model and observations for the main frequency range, which one finds by visual inspection of the above curves, is typical for all 32 stations in List B. A closer inspection shows that there is also satisfactory correspondence for the extremely high values of R , with no clearcut systematic difference, when all 32×12 distributions are considered.

However, the model seems to underestimate somewhat the risks of very dry events, approximately $R/E[R] \leq 0.15$, for a majority of months and stations in Southern Norway, whereas in Northern Norway this tendency is less obvious. The correction

transformation

$$F^* = 0.01 \cdot \left(1 - e^{-\frac{R}{E[R] \cdot 0.05}} \right) + 0.99 F(R/E[R]) \quad (61)$$

$$R^* = \frac{R}{0.99}$$

may be used to correct the tabulated values for $n = 30$, as a preliminary measure, until more is known about the reasons for the mentioned discrepancy.

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Note added in proof. According to observing practices at Norwegian meteorological stations, the numbers in

Table 6 express the probabilities that precipitation is falling within the time it takes to make the full set of observations, approx. a quarter of an hour. As compared with instantaneous precipitation probabilities and intensities, the numbers in Tables 6 and 8 will therefore be somewhat too large and small, respectively, and the numbers in Table 7 too large.

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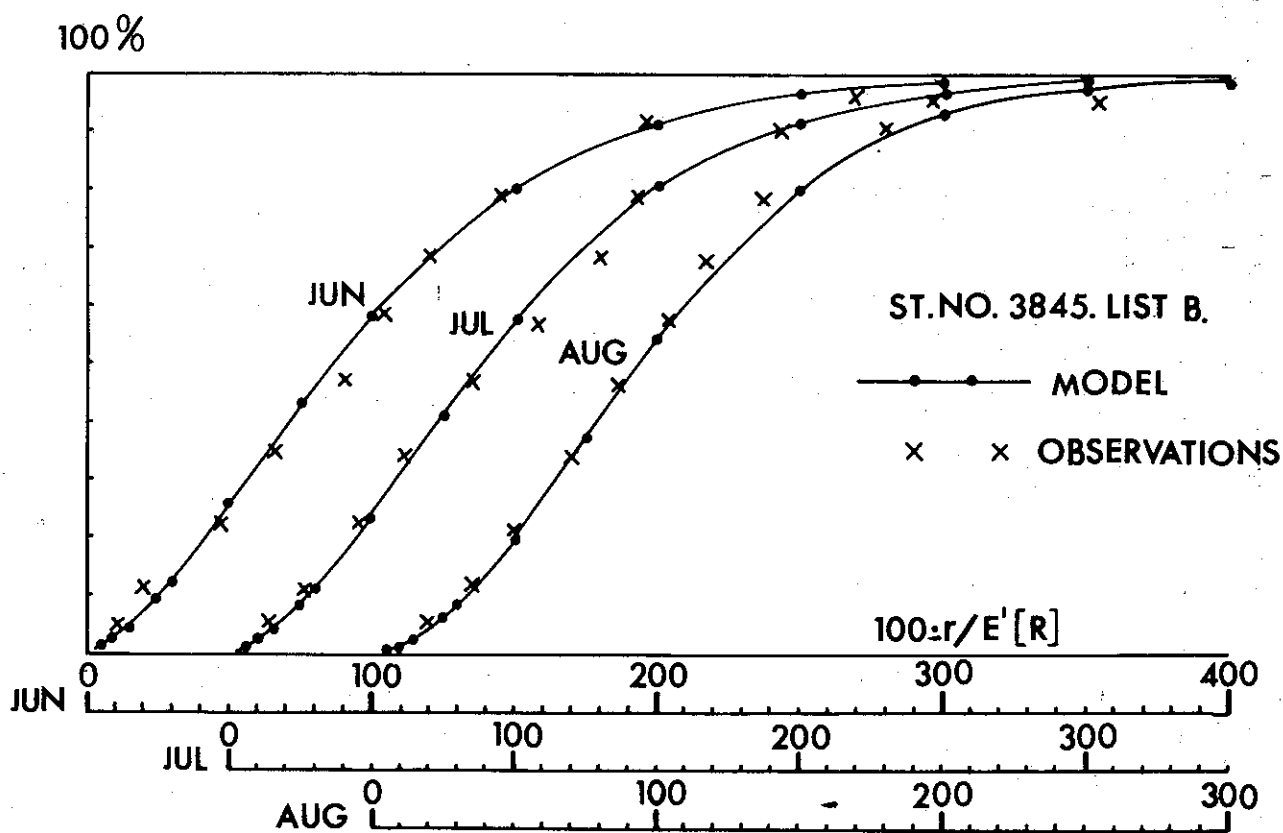
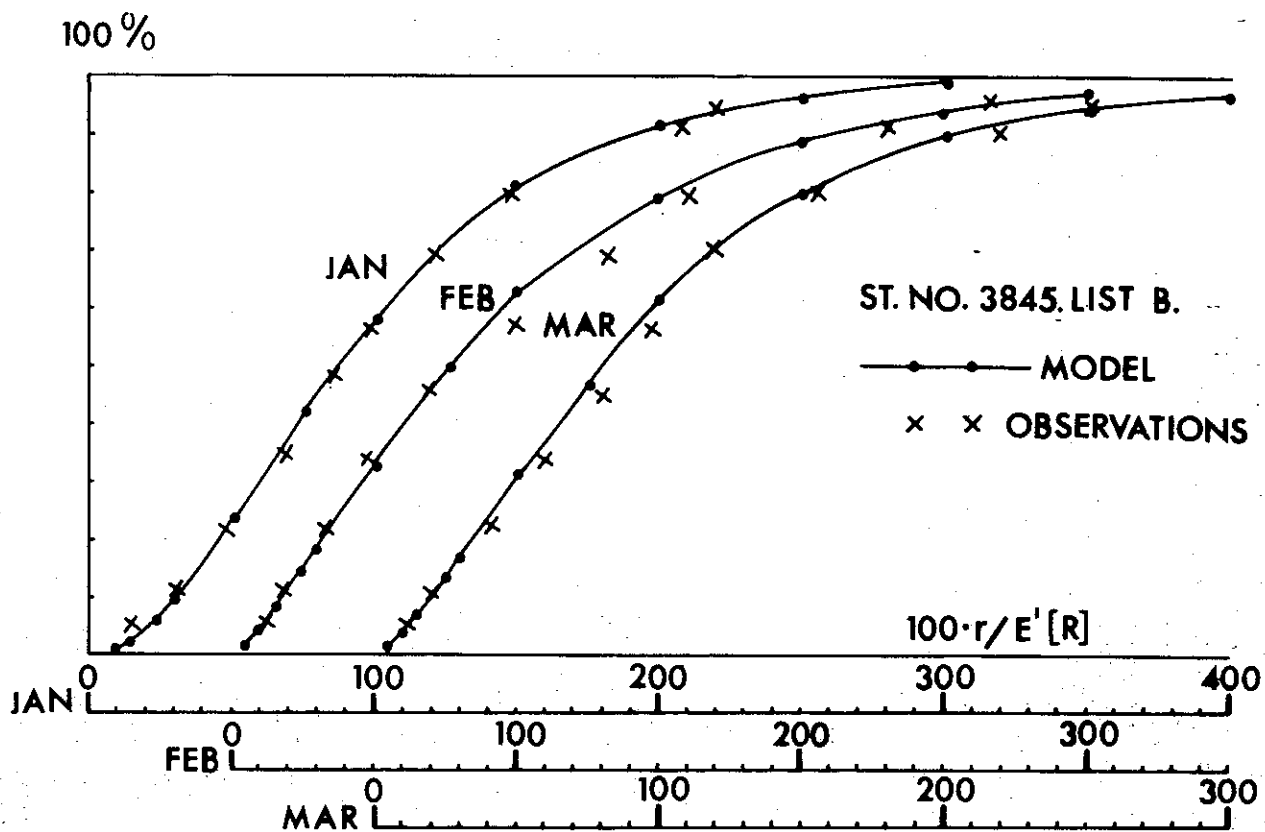


Fig. 2. These diagrams represent examples of probability distributions of 30 day precipitation sums as percentages of estimates of the corresponding expectancies. Dotted points represent values computed from model, while crosses are frequencies obtained by direct counting from observation series. The diagrams are drawn in order to give a visual impression of the typical goodness of the model in the main frequency range.

297 267 261 238 220 226 233 234 294 310 326 316

145 97 113 110 94 101 100 77 118 138 164 152

Station List A. A station number referred to in the text corresponds with its sequence number in the station list, which for convenience is arranged in groups of ten and ten stations, headed by the first and last numbers of each group. The first column contains the official archive numbers of the stations, the next three their geographical coordinates, latitudes, longitudes in degrees and minutes, and heights above sea level in meters, respectively.

0020	6120/1215/ 356	ST.NO 1 - 10	ST.NO 81 - 90	ST.NO 161 - 170	ST.NO 241 - 250
0070	6153/1203/ 672	3450 5900/0913/ 67	6537 6325/0804/ 30	9371 6901/2304/ 330	
0113	5900/1132/ 157	3586 5838/0909/ 4	6595 6351/0828/ 28	9390 6845/2332/ 382	
0140	5909/1134/ 114	3620 5824/0848/ 12	6610 6320/0939/ 300	9426 7041/2340/ 69	
0290	5959/1207/ 150	3656 5839/0838/ 169	6670 6250/1001/ 424	9450 7106/2400/ 13	
0307	5920/1054/ 34	3723 5902/0831/ 252	6671 6250/1001/ 441	9470 7104/2614/ 33	
0315	5919/1103/ 57	3814 5820/0831/ 6	6683 6237/1015/ 543	9489 7027/2513/ 12	
0341	5930/1117/ 141	3904 5812/0805/ 12	6817 6325/1026/ 113	9490 7027/2513/ 10	
0398	5942/1118/ 154	3910 5804/0803/ 9	6830 6312/1107/ 197	9535 7004/2459/ 5	
0478	6012/1105/ 202	3917 5810/0759/ 22	6886 6325/1027/ 127	9543 7004/2507/ 35	
		3969 5840/0748/ 212	6907 6319/1056/ 671	9640 7105/2814/ 8	
		ST.NO 91 - 100	ST.NO 171 - 180	ST.NO 251 - 260	
0487	6004/1116/ 247	3971 5840/0749/ 206	6910 6328/1056/ 12	9680 7024/2812/ 9	
0493	6006/1123/ 162	4014 5906/0730/ 443	6936 6325/1146/ 218	9708 6954/2625/ 112	
0565	6013/1201/ 175	4090 5938/0726/ 920	6995 6348/1112/ 76	9725 6928/2530/ 129	
0604	6037/1201/ 184	4111 5803/0727/ 138	7036 6340/1201/ 251	9735 6922/2426/ 286	
0701	6110/1127/ 240	4166 5815/0719/ 260	7085 6410/1229/ 195	9840 7042/3005/ 9	
0755	6122/1123/ 262	4177 5759/0703/ 13	7091 6415/1225/ 127	9855 7022/3106/ 14	
0825	6137/1054/ 303	4216 5807/0634/ 14	7155 6342/0936/ 9	9870 7005/3006/ 6	
0871	6153/1009/ 739	4280 5840/0642/ 57	7165 6351/0944/ 4	9915 6951/2925/ 8	
0885	6202/1047/ 485	4334 5826/0554/ 63	7199 6424/1027/ 18	9937 6944/2954/ 89	
0960	6218/1045/ 483	4350 5833/0621/ 196	7214 6413/1113/ 85	9953 6910/2915/ 54	
		ST.NO 101 - 110	ST.NO 181 - 190	ST.NO 261 - 265	
1040	6234/1123/ 628	4408 5839/0533/ 24	7285 6436/1216/ 22	9971 7431/1901/ 16	
1103	6009/1127/ 152	4432 5848/0538/ 14	7347 6427/1336/ 403	9972 7630/2504/ 6	
1150	6042/1052/ 264	4456 5853/0538/ 8	7349 6427/1343/ 462	9979 7804/1338/ 7	
1209	6044/1106/ 153	4464 5857/0544/ 72	7541 6448/1033/ 33	9986 7813/1535/ 37	
1218	6048/1112/ 200	4590 5911/0604/ 1	7560 6506/1142/ 47	9995 7056/0840/ 10	
1255	6046/1048/ 128	4651 5950/0659/1079	7632 6528/1213/ 5		
1264	6105/1029/ 271	4661 5939/0622/ 5	7650 6552/1211/ 6		
1266	6106/1029/ 226	4691 5929/0545/ 64	7719 6551/1312/ 4		
1354	6135/0947/ 249	4720 5909/0515/ 7	7742 6511/1325/ 339		
1355	6136/0946/ 241	4730 5918/0453/ 55	7765 6537/1402/ 380		
		ST.NO 111 - 120	ST.NO 191 - 200		
1367	6131/0923/ 865	4790 5951/0600/ 24	7795 6553/1318/ 634		
1431	6144/0933/ 285	4833 5955/0504/ 15	7940 6616/1359/ 31		
1460	6152/0906/ 371	4949 6019/0639/ 12	7948 6619/1410/ 51		
1469	6149/0858/ 746	4991 6034/0656/ 60	7980 6630/1514/ 549		
1536	6142/0817/ 674	5013 6013/0559/ 1	8010 6624/1237/ 10		
1554	6152/0827/ 378	5030 6024/0555/ 408	8070 6649/1359/ 39		
1572	6154/0752/ 712	5046 6016/0521/ 50	8145 6710/1501/ 793		
1655	6204/0908/ 643	5050 6018/0513/ 48	8162 6658/1518/ 26		
1660	6207/0917/ 952	5054 6023/0520/ 39	8180 6634/1521/ 680		
1661	6207/0917/ 972	5056 6024/0519/ 41	8208 6715/1523/ 14		

Station List B. Stations measuring 24 hr. precipitation amounts. Columns as for Station List A.

1715	5923/1047/	41	50
1729	5926/1036/	12	40
1785	5940/1047/	95	12
1870	5957/1043/	94	95
1895	5959/1041/	514	94
1940	5954/1037/	10	514
1948	5954/1031/	59	10
1971	5951/1026/	154	59
2036	6012/1019/	192	154
2124	6047/1012/	159	192
ST. NO 51 - 60			
2167	6056/1007/	485	60
2316	6055/0918/	634	485
2350	6107/0904/	525	634
2354	6114/0856/	822	525
2483	6033/0910/	369	822
2487	6034/0908/	165	369
2496	6043/0857/	542	165
2559	6031/0813/	810	542
2561	6032/0813/	768	810
2573	6031/0752/	988	768
ST. NO 61 - 70			
2584	6036/0730/1224	70	988
2590	6037/0725/1300	58	70
2648	5953/0954/	58	58
2689	5944/1012/	5	58
2723	5919/1031/	31	5
2735	5914/1017/	76	31
2741	5909/1027/	43	76
2745	5914/1021/	26	43
2747	5912/1016/	90	26
2750	5902/1032/	6	90
ST. NO 71 - 80			
2836	5940/0939/	171	6
2880	5954/0932/	288	171
2977	6025/0827/	871	288
3161	5951/0804/	948	871
3197	5951/0840/1828		948
3210	5923/0911/	26	
3242	5933/0825/	614	26
3243	5933/0825/	655	614
3306	5927/0800/	77	655
3412	5852/0936/	12	77
ST. NO 121 - 130			
5159	6039/0630/	130	12
5167	6041/0645/	560	130
5230	6050/0556/	104	560
5253	6045/0443/	20	104
5286	6102/0523/	39	20
5310	6110/0639/	51	39
5413	6104/0731/	36	51
5516	6130/0742/	27	36
5523	6131/0754/2062		27
5535	6126/0725/	484	
ST. NO 131 - 140			
5543	6140/0717/	324	484
5578	6111/0652/	53	324
5584	6126/0646/	10	53
5718	6128/0551/	41	10
5775	6134/0448/	10	41
5776	6134/0448/	9	10
5807	6147/0611/	51	9
5870	6156/0714/	201	51
5880	6156/0606/	71	201
5910	6202/0459/	38	71
ST. NO 141 - 150			
5958	6222/0600/	30	38
5961	6206/0535/	41	30
5971	6212/0608/	35	41
5980	6220/0516/	38	35
6020	6218/0648/	84	38
6050	6214/0725/	8	84
6065	6218/0714/	50	8
6083	6230/0641/	26	50
6099	6234/0607/	22	26
6104	6236/0619/	13	22
ST. NO 151 - 160			
6115	6237/0710/	49	13
6177	6214/0822/	621	49
6249	6252/0632/	8	621
6265	6258/0709/	26	8
6330	6224/0834/	869	26
6350	6233/0906/	195	869
6390	6218/0936/	885	195
6426	6307/0745/	48	885
6458	6250/0831/	3	48
6510	6312/0900/	9	3
ST. NO 201 - 210			
8229	6716/1422/	11	9
8240	6724/1354/	14	11
8312	6750/1447/	6	14
8370	6803/1602/	60	6
8479	6828/1730/	32	60
8490	6826/1805/	512	32
8524	6819/1539/	20	512
8538	6809/1439/	11	20
8578	6753/1303/	31	11
8590	6730/1204/	8	31
ST. NO 211 - 220			
8595	6725/1153/	18	8
8626	6821/1358/	10	18
8652	6839/1517/	23	10
8676	6838/1428/	12	23
8710	6919/1607/	5	12
8711	6918/1609/	10	5
8735	6846/1612/	36	10
8742	6856/1641/	45	36
8779	6835/1635/	7	45
8800	6845/1749/	20	7
ST. NO 221 - 230			
8868	6933/1755/	9	20
8890	6921/1805/	12	9
8935	6903/1833/	76	12
8980	6901/1917/	78	76
8995	6847/1943/	228	78
9008	6920/1854/	22	228
9020	6915/1914/	22	22
9028	6938/1801/	2	22
9045	6939/1856/	100	2
9049	6941/1855/	8	100
ST. NO 231 - 240			
9080	7015/1930/	21	8
9126	6934/2014/	3	21
9133	6915/1959/	37	3
9135	6923/2018/	46	37
9175	6945/2102/	4	46
9235	6950/2153/	6	4
9270	7020/2128/	10	6
9314	6959/2322/	3	10
9330	6935/2332/	374	3
9370	6900/2302/	306	374
ST. NO 201 - 210			
0060	6151/1151/	696	306
0165	5918/1140/	113	696
0535	6023/1134/	147	113
0655	6103/1153/	513	147
1010	6231/1101/	788	513
1252	6047/1058/	205	788
1835	5956/1045/	83	205
2012	6007/1024/	442	83
2188	6102/0931/	679	442
2780	5912/0958/	31	679
3490	5916/0846/	464	31
3845	5831/0821/	85	464
3922	5813/0754/	151	85
4289	5849/0643/	475	151
4448	5841/0559/	263	475
4480	5849/0555/	230	263
4645	5950/0650/	393	230
5035	6028/0554/	370	393
5056	6024/0519/	41	370
5147	6039/0614/	319	41
5711	6126/0513/	162	319
5896	6200/0639/	340	162
6522	6316/0900/	133	340
7210	6415/1112/	86	133
7380	6441/1339/	376	86
7785	6522/1416/	500	376
8020	6623/1312/	110	500
8190	6708/1605/	142	110
8350	6748/1559/	76	142
8685	6849/1448/	3	76
9175	6945/2102/	4	3
9330	6935/2332/	374	4
9560	7019/2533/	10	374

Table 6

Estimates of precipitation probabilities based upon observations on the main observing hours, 00, 06, 12, 18 GMT or 06, 12, 18 GMT, and averaged for calendar months Jan., . . . , Dec. The numbers are arranged horizontally according to months, and vertically in 10 and 10 stations according to Station List B. Unit: 1/1000.

J	F	M	A	M	J	J	J	A	S	O	N	D	ST. NO											
													10	10	10	10	10	10	10	10	10	10	10	10
228	188	154	190	159	131	172	147	152	198	281	246	269	204	152	132	109	109	79	105	106	117	173	243	221
184	166	164	164	135	139	151	129	125	149	187	188	213	174	134	110	103	78	83	90	111	157	206	190	
217	159	145	133	101	71	88	71	111	157	216	183	255	196	153	136	116	97	98	107	145	182	235	230	
190	185	125	178	140	117	165	194	167	226	271	234	293	210	171	141	122	74	99	81	136	207	238	212	
165	137	104	135	122	106	128	110	132	167	202	166	265	191	158	146	136	113	140	135	165	199	258	224	
175	133	99	125	126	100	113	102	133	179	233	179	251	183	156	129	108	80	95	97	117	167	226	227	
167	126	104	99	86	87	86	80	104	147	171	135	309	255	199	158	140	103	107	105	145	192	280	284	
150	94	92	76	82	71	86	58	95	141	145	118	254	209	172	141	121	95	106	108	146	183	255	251	
223	184	145	160	140	137	143	140	187	206	260	235	221	170	149	137	110	77	95	102	114	147	228	210	
286	230	165	164	139	109	135	119	148	202	261	257	410	298	226	173	132	97	132	94	163	213	307	305	
													ST. NO 91											
250	185	154	159	112	98	113	110	145	192	244	230	243	163	139	136	151	103	127	144	144	196	242	253	
182	156	118	123	100	90	107	101	107	153	190	167	437	321	290	239	233	211	237	226	356	312	407	394	
205	179	123	136	112	83	101	100	112	151	219	213	217	113	132	103	44	72	81	73	167	140	169	188	
154	111	81	103	86	86	98	84	97	114	148	143	223	171	168	141	119	93	112	119	138	199	238	229	
144	101	88	102	109	97	132	101	117	129	170	148	473	364	325	233	228	169	183	161	275	304	400	397	
261	177	147	127	135	103	113	94	150	197	213	220	266	213	185	151	137	113	126	135	186	213	294	263	
183	165	112	129	122	88	122	129	107	152	233	233	229	185	160	146	128	114	121	129	161	206	248	224	
209	174	173	184	134	120	141	109	138	161	222	224	212	167	196	167	177	160	171	175	190	249	254	247	
165	144	131	146	133	109	142	118	124	139	202	182	229	174	152	153	145	113	123	113	209	223	270	232	
183	135	111	111	104	106	139	103	97	123	178	195	452	368	328	285	239	211	228	177	353	328	386	387	
													ST. NO 101											
243	225	199	218	146	143	165	135	177	196	252	263	217	174	150	165	140	148	143	147	171	212	238	224	
167	159	89	117	102	94	120	113	122	137	228	212	122	90	94	92	92	86	83	97	92	140	153	137	
185	151	122	125	97	96	123	101	117	150	191	194	223	193	182	175	163	165	157	156	201	224	264	230	
122	103	81	100	98	81	117	100	95	112	144	129	206	169	150	143	154	138	135	147	190	209	261	236	
161	150	102	142	140	113	154	127	136	156	244	211	238	190	198	193	177	203	189	195	265	268	307	279	
196	145	116	129	108	103	126	108	123	143	199	196	423	333	326	317	198	242	270	202	400	419	454	378	
206	145	161	117	106	143	132	90	137	161	197	173	311	257	278	230	216	219	213	226	310	341	380	356	
174	153	104	136	141	108	146	129	139	175	228	207	261	203	214	188	152	170	179	147	245	235	318	290	
167	90	91	79	56	83	102	67	86	83	139	120	182	156	130	132	125	122	118	129	186	188	220	192	
151	156	97	103	132	139	191	182	133	135	206	246	217	190	163	145	129	126	142	136	190	210	248	231	
													ST. NO 111											
135	61	79	97	53	70	106	67	70	73	128	102	245	215	226	211	189	204	228	222	243	297	248	260	
90	51	57	48	60	117	129	68	100	86	111	82	194	171	158	128	123	141	126	133	196	199	249	212	
171	113	90	82	98	98	111	90	108	124	180	174	256	193	215	184	153	179	181	151	279	274	298	315	
268	174	145	123	90	130	135	87	123	148	213	245	213	173	223	153	126	142	176	132	250	223	270	316	
121	112	126	81	121	125	129	112	146	154	144	134	269	233	243	208	182	187	192	183	283	272	320	323	
168	89	110	90	90	117	184	104	193	139	177	177	293	260	279	228	187	238	224	220	311	309	342	346	
276	208	191	119	108	124	129	104	205	196	253	305	207	182	204	178	152	184	163	146	237	220	253	255	
183	166	149	142	121	150	157	140	149	174	198	202	257	233	213	200	168	184	188	177	225	245	274	275	
181	101	158	127	142	141	168	130	120	167	205	230	268	239	233	212	182	200	205	185	259	275	276	276	

236	194	153	130	114	ST.NO 41	50	120	171	223	195	290	236	234	189	164	172	191	151	290	300	369	330
197	132	126	135	110	81	85	92	122	153	202	145	243	237	192	177	243	247	238	329	303	314	287
163	125	115	114	90	75	105	80	99	133	165	132	285	284	229	187	233	229	198	311	336	336	340
254	199	161	161	136	109	129	111	145	168	241	220	294	255	238	176	171	178	189	267	293	326	313
192	157	148	149	131	100	136	105	132	158	219	195	282	247	243	177	226	230	198	291	303	314	317
255	206	146	141	116	96	121	105	123	173	223	210	126	116	102	63	88	98	105	170	171	162	156
319	230	210	160	123	113	145	119	147	206	223	194	158	125	143	136	162	173	163	226	202	193	203
192	177	121	135	115	90	126	97	122	152	212	174	182	143	143	108	120	125	112	210	194	201	220
159	119	97	113	101	90	125	110	118	135	207	178	318	290	279	254	238	248	240	286	320	378	346
212	165	131	130	118	117	129	114	130	152	218	208	213	210	220	142	140	157	137	208	266	227	238
191	157	137	144	122	124	129	108	131	150	184	168	179	124	167	107	131	137	117	210	208	210	224
137	114	114	110	67	75	93	88	106	115	142	135	212	199	189	138	123	152	154	232	231	232	238
215	165	156	137	142	116	122	118	157	161	212	205	303	262	268	197	173	209	213	195	313	345	356
304	212	222	154	161	157	153	121	196	218	233	254	304	256	309	200	182	217	219	211	344	366	368
174	149	123	137	135	120	142	119	147	123	253	216	452	366	347	269	199	244	252	242	394	454	519
156	122	92	98	100	102	123	98	113	121	186	173	254	242	235	218	167	185	182	185	218	249	267
151	113	99	106	101	104	131	105	125	109	142	137	214	148	173	129	103	138	135	140	232	220	255
298	218	196	174	127	113	164	112	183	187	267	256	191	179	164	148	112	136	158	129	217	216	203
119	135	97	97	90	87	126	118	111	129	193	154	320	304	286	219	214	217	235	203	292	316	271
287	228	231	192	168	151	158	151	221	239	305	311	228	187	178	158	121	138	155	139	188	231	228
426	379	361	273	197	187	216	210	353	335	473	452	315	295	336	221	194	245	299	202	315	345	361
287	286	301	215	208	261	258	251	330	355	325	329	294	259	261	243	213	233	258	220	373	340	334
247	191	140	143	126	107	119	108	145	182	232	218	188	201	209	144	110	154	171	151	246	233	236
297	187	139	133	113	74	103	74	127	190	244	148	272	246	236	189	170	202	214	190	291	290	311
215	161	122	156	125	115	106	113	133	181	240	210	145	175	194	98	85	95	122	76	160	164	196
237	196	120	131	116	93	101	113	112	171	264	228	210	201	186	180	129	169	170	168	236	234	226
231	165	126	117	91	67	94	65	114	159	192	149	173	192	207	119	90	114	145	108	171	198	179
218	184	124	122	94	73	91	86	100	152	222	181	238	241	229	214	161	184	216	175	305	280	282
265	198	145	128	110	92	98	92	119	172	243	215	203	191	169	163	134	173	176	154	233	227	223
195	165	119	114	95	75	83	91	106	158	193	160	142	156	155	147	123	107	161	129	222	188	228
244	194	139	143	122	97	135	109	142	179	238	205	196	188	183	158	141	165	170	151	220	230	182
244	186	147	159	133	109	142	121	141	175	246	229	86	167	108	111	65	111	140	108	175	121	125
108	83	78	76	65	68	92	72	92	106	146	123	213	202	192	174	117	151	183	142	212	232	223
325	250	211	205	207	161	175	145	242	242	319	278	224	209	206	190	144	197	202	163	250	273	249
195	175	136	137	179	146	168	163	159	161	247	205	249	261	222	177	109	119	129	107	173	196	248
194	154	112	120	105	93	127	112	135	174	203	167	171	174	169	152	89	107	119	102	143	156	148
351	266	186	203	181	120	147	122	170	171	263	254	210	179	145	167	129	100	194	129	103	156	148
351	262	190	203	181	117	147	125	173	171	256	258	212	222	185	202	154	184	194	146	254	267	232
272	213	160	132	129	95	133	115	157	179	244	228	220	239	223	183	167	206	202	177	231	218	183
240	176	126	128	111	90	105	102	125	182	243	203	259	234	236	214	170	212	208	191	283	278	254

222	220	194	183	140	178	206	163	214	278	229	227
256	256	252	198	145	227	236	200	290	339	269	294
119	126	110	109	72	103	111	95	123	171	129	120
169	137	167	154	131	144	140	132	179	228	160	148
175	183	173	158	114	137	144	132	163	219	167	176
344	354	366	325	265	290	237	233	300	348	294	273
215	213	247	193	145	187	197	134	204	318	256	219
247	230	208	178	132	151	149	143	190	272	241	270
331	316	311	250	173	222	213	192	300	369	338	362
174	153	159	136	82	113	114	107	126	179	142	177
ST. NO 201 - 210											
213	196	167	153	95	138	150	120	165	217	194	226
177	171	164	143	94	125	129	104	143	196	188	181
331	341	318	276	220	241	252	240	313	393	349	343
228	234	219	196	151	171	188	160	221	275	253	248
207	199	194	167	125	152	138	106	187	251	223	189
216	235	217	215	158	181	194	180	215	274	242	229
234	243	216	192	143	127	157	159	219	240	227	250
165	199	181	183	113	170	219	206	177	226	193	204
255	272	253	213	195	225	237	229	270	296	241	261
153	210	200	180	138	138	151	155	176	234	196	196
ST. NO 211 - 220											
188	221	185	183	144	170	185	185	227	258	258	242
240	236	250	179	144	154	167	134	215	253	235	247
247	263	217	207	168	181	194	175	209	275	223	245
183	225	187	178	157	158	181	173	223	248	223	202
104	120	77	81	101	113	139	121	113	118	110	90
226	261	258	243	155	177	194	206	220	316	245	232
355	316	161	167	97	167	161	194	200	226	300	226
202	187	188	155	147	191	174	165	203	223	180	152
295	285	276	264	208	232	224	230	290	320	286	272
254	291	256	231	185	194	193	217	256	334	294	282
ST. NO 221 - 230											
194	193	181	149	136	171	233	163	167	217	207	195
175	209	207	164	126	128	151	158	167	243	186	185
212	213	226	226	154	167	189	208	231	278	240	228
117	115	89	58	34	58	39	67	80	103	82	75
287	315	278	249	237	251	241	244	276	277	264	264
152	172	159	148	120	140	136	166	149	200	180	144
258	259	241	187	169	181	151	195	226	247	203	220
183	177	169	123	103	107	91	110	138	162	188	176
236	215	208	250	200	180	149	198	221	258	261	241
124	94	97	98	108	108	102	124	127	122	136	129
ST. NO 231 - 240											

89	86	84	68	ST.NO 41	50	51	71	89	88	62	ST.NO 121	130	78	108	101	81
91	82	73	70	54	48	47	84	87	76	75	61	58	115	136	130	130
79	75	74	71	58	64	56	67	80	77	99	95	87	108	124	104	121
116	109	88	71	52	57	60	68	107	120	75	80	77	74	71	79	83
83	79	94	70	62	50	50	57	95	105	81	73	68	116	107	102	115
99	100	88	76	64	53	49	71	83	101	42	46	45	58	61	59	70
115	130	89	69	56	54	54	87	88	79	43	51	55	70	60	80	77
109	129	93	76	66	66	60	78	104	97	46	42	47	62	56	71	87
119	74	62	58	57	54	75	80	105	108	78	89	77	76	80	91	88
112	90	85	78	68	59	58	81	100	110	100	76	59	72	99	103	106
100	97	90	80	ST.NO 51	60	65	83	84	122	59	ST.NO 131	140	58	60	77	88
87	93	71	64	56	51	52	90	76	77	53	58	75	70	84	81	87
101	79	78	70	42	50	41	78	95	106	65	55	50	101	119	121	164
106	118	131	87	66	52	51	88	98	113	80	79	83	124	133	128	159
86	93	61	44	76	70	73	58	112	100	86	100	97	119	157	163	136
101	83	65	57	53	48	49	58	102	109	71	77	86	119	60	71	164
83	95	68	57	63	48	50	69	80	82	72	64	62	65	60	74	84
145	117	91	95	57	44	57	52	69	101	60	50	75	82	72	74	80
69	109	64	44	65	57	71	109	98	101	50	59	52	59	69	80	80
108	87	79	68	57	59	56	65	88	82	83	83	89	79	80	106	139
124	150	160	99	ST.NO 61	70	91	94	94	102	53	52	53	56	58	53	61
109	90	132	91	96	99	90	131	144	155	91	ST.NO 141	150	98	116	134	130
132	127	99	77	86	131	100	107	113	113	79	114	136	88	87	87	75
114	127	69	88	74	53	53	103	127	147	71	72	76	82	84	88	86
104	82	68	65	68	46	56	113	94	99	71	60	61	82	67	59	66
106	107	94	67	77	64	47	88	94	117	53	60	72	68	67	65	58
84	97	73	68	71	46	56	84	104	101	37	55	48	55	75	65	58
104	110	94	65	53	42	47	81	68	68	89	69	64	78	87	99	117
104	96	92	63	66	63	50	71	105	86	70	61	54	54	71	78	82
80	87	78	64	64	55	59	75	105	85	44	55	50	101	84	73	84
145	108	90	74	ST.NO 71	80	52	66	70	77	55	69	37	51	56	52	47
122	125	101	69	61	50	53	90	133	104	68	ST.NO 151	160	66	67	72	72
63	45	49	37	68	54	63	110	118	113	36	47	60	78	54	38	61
108	79	99	81	40	47	57	70	63	65	49	36	31	72	70	54	66
70	74	56	50	75	71	71	99	99	108	59	65	68	66	74	68	80
105	88	72	62	59	58	62	70	80	79	70	44	39	69	97	93	118
252	118	83	121	61	60	57	82	109	93	47	49	60	46	65	66	66
245	122	88	117	105	63	56	57	101	119	71	82	48	62	35	66	66
147	140	119	73	73	61	54	57	105	121	110	54	82	82	60	66	64
103	97	92	75	71	58	58	85	104	103	58	68	68	63	74	90	98
				62	67	62	81	97	90	70	69	69	68	67	67	85
				67	63	52	81	97	90	44	58	48	57	58	50	41
				66	50	80	66	70	77	55	48	51	51	58	50	41
				67	60	74	90	133	104	68	47	60	66	67	72	72
				66	63	60	66	110	113	36	60	51	78	54	38	61
				67	46	56	70	63	65	49	36	31	72	70	54	66
				68	71	87	99	99	108	59	65	68	66	74	68	80
				62	62	56	70	80	79	70	44	39	69	97	93	118
				57	60	79	82	109	93	47	49	60	46	65	66	66
				56	56	57	57	101	119	71	82	48	62	35	66	66
				54	54	54	57	101	119	110	54	82	82	60	66	64
				58	61	58	57	105	121	58	68	68	63	60	66	64
				58	58	58	71	85	104	78	68	68	63	74	90	98
				63	62	63	85	104	103	78	68	68	63	63	67	85
				67	67	63	81	97	90	44	58	48	57	58	50	41

64	63	58	68	ST.NO 201	59	55	64	48	55	72	60	60	60
65	88	69	77	55	48	67	59	48	75	70	71	76	76
56	46	45	53	70	54	48	54	59	45	72	50	39	58
78	53	57	76	40	48	59	68	45	62	65	74	58	62
59	72	53	67	47	46	45	62	45	45	77	60	62	106
134	121	172	107	88	66	66	61	61	98	117	112	106	67
75	62	64	85	54	55	55	65	65	45	75	60	67	64
57	55	52	59	54	54	49	49	49	55	74	55	64	84
76	78	91	76	66	104	80	80	80	82	108	86	84	61
47	41	54	53	48	48	60	41	41	43	52	36	61	48
53	41	54	54	ST.NO 211	49	53	57	40	52	59	44	48	52
56	49	46	60	47	53	51	47	47	45	66	54	52	147
127	141	154	125	106	96	101	99	99	116	154	129	147	58
61	54	66	64	55	59	62	47	47	61	70	50	58	54
53	58	59	62	58	60	50	53	53	65	62	56	54	47
53	58	60	61	55	56	58	59	59	54	64	51	47	66
82	77	103	61	48	52	54	52	52	67	74	60	48	66
57	64	63	50	63	61	56	48	48	47	83	57	48	91
82	67	81	74	75	86	67	78	78	75	110	78	91	74
74	82	76	65	68	50	63	49	49	67	89	81	74	53
52	56	59	62	ST.NO 221	47	47	46	46	42	53	53	53	74
83	92	89	88	63	56	71	55	55	54	71	76	83	95
83	85	72	74	62	55	59	58	58	56	76	67	83	34
93	99	79	75	59	64	73	69	69	79	86	83	95	81
59	49	38	42	53	54	74	70	70	64	63	50	34	81
84	104	89	115	79	55	60	58	58	57	86	87	81	51
112	84	93	39	82	110	0	28	28	27	75	106	51	37
42	50	38	60	54	62	63	53	53	64	53	57	37	84
92	100	73	102	71	61	76	70	70	67	102	85	84	87
78	95	98	87	72	71	64	72	72	64	94	78	87	50
52	45	51	48	ST.NO 231	59	92	63	63	47	61	51	50	92
82	97	107	125	62	56	63	57	57	74	72	90	92	91
118	108	103	125	67	82	77	59	59	70	91	98	91	65
66	72	53	45	34	38	28	41	41	51	64	77	65	132
130	126	128	95	92	66	73	77	77	79	100	101	132	55
64	69	64	57	64	53	57	69	69	60	84	66	55	106
134	107	115	80	87	91	81	102	102	75	107	98	106	50
50	53	68	46	45	38	40	46	46	35	48	50	50	72
89	60	48	74	75	62	54	54	54	70	59	76	72	60
56	44	64	53	61	56	56	47	47	55	62	58	60	58

Table 8

Estimates of the expected values of precipitation intensities greater than zero, averaged for calendar months. Unit: 0.01 mm hr⁻¹. Month and station arrangements as in Table 6.

J	F	M	A	M	J	J	A	S	O	N	D	ST.NO	81	90	ST.NO	81	90	ST.NO	81	90	ST.NO	81	90	ST.NO	81	90
19	23	25	31	53	88	82	81	79	59	34	28	75	114	117	84	119	112	75	114	117	84	119	112	75	114	117
20	17	19	23	38	61	81	76	65	38	31	25	78	87	113	50	49	46	87	87	113	50	49	46	87	87	113
39	37	49	52	77	115	116	114	127	82	67	50	64	74	87	101	98	103	64	74	87	101	98	103	64	74	87
47	36	32	39	59	86	77	87	63	67	47	44	85	126	121	60	61	117	85	126	121	60	61	117	85	126	121
37	35	40	45	64	82	90	100	86	59	52	41	71	84	94	72	72	102	71	84	94	72	72	102	71	84	94
43	43	46	46	86	86	82	123	86	83	60	49	98	120	124	81	94	117	98	120	124	81	94	117	98	120	124
49	51	55	65	94	108	121	140	119	91	79	65	81	124	154	81	94	117	81	124	154	81	94	117	81	94	117
53	70	64	82	86	133	112	150	149	92	85	69	73	86	99	60	73	86	73	86	99	60	73	86	73	86	99
30	33	38	44	66	76	87	84	78	73	48	31	106	131	142	72	106	131	106	131	142	72	106	131	106	131	142
29	29	36	43	59	90	96	101	87	66	53	37	77	98	90	51	77	98	77	98	90	51	77	98	77	98	90
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31	36	43	48	78	94	117	111	97	70	54	38	79	101	129	44	79	101	79	101	129	44	79	101	129	44	79
48	50	50	59	73	101	96	94	89	77	59	54	61	72	72	42	61	72	61	72	72	42	61	72	72	42	61
28	32	28	34	65	89	118	95	74	49	42	32	67	73	91	50	67	73	67	73	91	50	67	73	91	50	67
22	20	27	34	51	86	107	80	87	58	34	24	72	75	94	50	72	75	72	75	94	50	72	75	94	50	72
19	17	17	18	40	75	91	82	57	44	27	21	82	96	102	51	82	96	82	96	102	51	82	96	102	51	82
15	17	19	22	43	88	85	71	65	47	26	19	62	76	98	77	65	76	62	76	98	77	65	76	98	77	65
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17	18	20	15	25	53	57	30	36	27	20	20	62	70	76	65	70	76	62	70	76	65	70	76	65	70	76
34	32	43	45	75	98	101	98	89	85	45	35	97	99	133	58	99	133	97	99	133	58	99	133	97	99	133
27	24	27	34	54	82	93	93	74	55	43	28	40	56	63	46	56	63	40	56	63	46	56	63	40	56	63
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25	26	29	35	52	80	83	88	69	59	39	25	32	50	48	52	50	48	32	50	48	52	50	48	32	50	48
36	35	39	44	59	70	86	89	72	52	46	36	63	82	77	63	82	77	63	82	77	63	82	77	63	82	77
35	30	37	43	58	78	85	95	75	69	48	41	90	95	112	73	95	112	90	95	112	73	95	112	90	95	112
22	19	22	25	55	79	90	69	65	50	33	23	80	74	90	73	74	90	80	74	90	73	74	90	73	74	90
19	15	19	16	34	49	51	50	53	47	23	17	58	66	79	65	66	79	58	66	79	65	66	79	65	66	79
40	45	43	33	92	124	107	40	90	70	50	40	ST.NO	111	120	ST.NO	111	120	ST.NO	111	120	ST.NO	111	120	ST.NO	111	120
29	23	28	27	63	77	71	70	57	44	37	31	92	143	124	92	143	124	92	143	124	92	143	124	92	143	124
16	20	18	15	35	60	65	58	50	39	20	20	70	76	88	75	76	88	70	76	88	75	76	88	70	76	88
19	20	20	15	37	64	80	60	50	33	30	24	59	41	52	41	44	52	59	41	52	41	44	52	41	44	52
42	32	36	27	37	50	55	63	52	57	36	41	78	83	107	52	55	72	78	83	107	52	55	72	78	83	107
17	19	10	8	19	38	39	34	17	23	31	23	105	94	115	87	83	109	105	94	115	87	83	109	105	94	115
25	22	28	12	28	37	45	34	38	41	34	34	116	119	112	87	119	112	116	119	112	87	119	112	116	119	112
20	21	19	12	30	47	51	46	36	33	24	24	105	96	116	89	96	116	105	96	116	89	96	116	105	96	116
24	21	22	19	30	47	51	59	45	33	24	22	71	69	94	63	69	94	71	69	94	63	69	94	71	69	94

52	44	52	39	ST.NO 201 -	210	81	75	60	57
40	30	34	29	44	50	50	51	44	36
87	66	75	65	31	47	126	118	99	94
68	67	53	49	78	91	85	91	68	74
57	51	41	41	62	80	75	74	52	54
34	29	22	21	63	75	44	51	32	33
54	53	57	49	37	48	79	73	60	69
34	33	35	34	47	68	83	61	51	43
68	59	57	56	43	52	61	57	75	65
55	51	53	42	48	65	79	92	72	63
				52	63	82	73		
				45	63	47			
				ST.NO 211 -	220				
37	32	37	28	38	32	51	47	44	39
85	76	77	64	62	59	107	103	99	100
54	54	47	40	35	39	64	70	62	54
47	47	46	42	36	45	74	73	58	56
50	48	48	42	44	47	68	65	50	55
41	40	41	34	37	37	61	60	55	51
44	44	41	33	29	44	54	58	56	46
64	52	61	40	39	41	69	64	58	62
36	35	31	26	24	28	41	45	40	39
59	51	40	37	36	47	66	62	59	56
				ST.NO 221 -	230				
61	64	53	40	42	43	70	65	65	60
47	46	38	36	47	47	61	61	51	49
35	34	27	20	19	32	45	37	36	35
30	37	28	25	22	41	41	46	36	39
21	19	15	12	20	48	37	30	21	21
38	46	35	31	20	55	52	42	40	42
57	49	47	22	46	58	57	78	77	58
61	54	49	39	55	76	70	76	63	65
43	44	40	31	49	45	53	53	49	45
36	44	38	31	30	37	54	50	46	41
				ST.NO 231 -	240				
52	46	45	38	47	42	68	58	51	51
44	51	41	36	23	45	45	46	55	50
28	34	19	19	21	41	35	37	31	32
37	32	29	22	55	56	54	48	44	42
27	25	24	17	27	34	34	36	31	28
25	27	30	18	21	35	42	34	26	30
39	41	40	36	48	59	60	59	52	42
20	20	22	19	50	86	40	60	20	23
15	17	15	11	24	63	29	31	17	16
13	13	10	12	60	85	41	20	17	15

2780
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 5 74 253 501 732 887 964 992 999 *** ***
 9 109 325 582 792 918 976 995 999 *** ***
 9 116 347 614 820 935 983 997 *** ***
 5 84 295 572 803 933 984 998 *** ***
 2 41 187 441 706 888 971 996 *** ***
 1 20 110 310 576 807 939 988 999 *** ***
 1 17 93 268 516 753 910 979 997 *** ***
 1 25 113 291 527 749 900 973 996 *** ***
 1 27 118 294 524 741 891 968 994 *** ***
 1 19 98 265 496 723 883 965 994 *** ***
 1 18 99 276 517 746 900 973 996 *** ***

3490
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 1 24 99 242 439 646 817 927 981 998 ***
 2 34 137 314 531 731 876 956 990 999 ***
 2 35 151 357 596 797 921 978 996 ***
 1 28 130 329 579 793 925 981 997 ***
 1 14 80 232 459 695 870 962 993 ***
 0 6 41 145 333 570 785 924 984 999 ***
 0 5 33 118 285 512 738 898 976 998 ***
 0 7 42 133 298 516 733 892 973 997 ***
 0 7 40 129 291 507 723 884 969 996 ***
 0 4 27 101 252 470 699 873 966 996 ***
 0 3 25 99 252 473 703 877 967 996 ***

3845
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 1 20 93 244 454 671 842 943 986 998 ***
 1 25 116 295 526 742 892 967 994 ***
 1 27 128 321 560 772 910 975 996 ***
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 0 11 59 182 385 623 824 943 990 999 ***
 0 7 37 122 284 509 737 903 980 999 ***
 0 6 34 108 248 446 663 845 954 995 ***
 0 4 25 84 200 374 581 775 913 983 998 ***
 0 1 8 38 117 268 483 713 890 979 998 ***
 0 0 5 27 96 245 476 726 909 987 999 ***

655
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 1 24 129 336 589 799 925 980 997 *** ***
 1 36 173 411 662 848 948 987 998 *** ***
 2 41 185 430 686 868 959 991 999 *** ***
 1 33 149 364 617 822 939 986 998 *** ***
 1 16 84 237 459 689 862 957 992 999 *** ***
 0 6 43 147 335 571 785 923 983 999 *** ***
 0 6 42 145 335 576 793 930 986 999 *** ***
 0 12 68 207 426 665 852 954 992 999 *** ***
 0 13 79 234 462 694 866 957 991 999 *** ***
 0 8 58 193 413 655 844 950 990 999 *** ***
 0 6 49 174 392 644 845 954 992 999 *** ***

1010
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 0 2 24 104 278 524 762 917 983 999 ***
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 0 10 66 208 433 677 862 960 993 ***
 0 4 32 122 302 545 774 923 985 999 ***
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 0 1 9 48 162 373 637 855 967 997 ***
 0 0 7 43 154 370 640 861 970 998 ***
 0 0 4 25 103 279 544 801 952 996 ***
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1252
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 1 21 110 302 557 784 924 983 998 *** ***
 1 32 144 351 598 804 929 982 997 *** ***
 1 30 139 341 585 792 921 979 997 *** ***
 1 24 122 322 573 791 925 982 997 *** ***

4480	0	5	26	79	180	334	526	722	879	971	996
	0	8	39	113	247	432	637	816	935	989	999
	0	8	41	128	287	502	720	885	971	997	***
	0	9	48	150	330	558	772	916	982	999	***
	1	12	62	178	365	587	785	918	980	998	***
	0	10	52	155	328	544	750	898	973	997	***
	0	3	22	81	209	409	640	837	954	995	***
	0	1	8	36	113	263	484	723	903	986	999
	0	1	10	37	103	229	419	643	844	967	997
	0	3	15	49	120	243	419	625	820	955	995
	0	3	15	47	118	241	418	627	823	957	995
	0	3	15	50	125	256	440	650	839	961	996

4645	1	13	51	133	267	445	638	809	928	987	999
	2	27	91	209	376	565	743	878	959	993	999
	2	30	113	266	468	676	840	941	986	999	***
	2	35	138	323	550	757	898	969	994	***	***
	3	42	159	351	575	770	900	968	993	999	***
	2	27	112	273	488	700	860	951	989	999	***
	0	7	41	133	304	531	751	905	978	998	***
	0	3	18	68	183	372	605	815	946	994	***
	0	6	29	87	199	367	571	767	912	984	998
	1	9	39	109	233	407	605	787	918	984	998
	0	5	26	87	208	393	609	804	934	990	999
	0	4	24	81	197	377	594	795	931	990	999

5035	0	5	25	74	167	311	496	691	857	964	996
	0	8	36	102	221	391	589	775	912	983	999
	1	10	46	133	285	488	699	866	961	995	***
	1	12	57	166	344	562	765	906	976	997	***
	0	12	58	163	332	540	739	885	966	995	***
	0	7	37	111	248	440	651	832	946	993	***
	0	3	17	61	157	321	538	758	918	990	999
	0	2	10	39	111	246	446	676	871	977	998
	0	2	11	39	108	233	419	636	833	960	995
	0	2	12	42	113	241	427	643	837	961	995
	0	2	12	42	114	246	439	661	854	969	997
	0	3	15	50	128	263	453	666	851	966	996

5896	1	12	46	116	230	387	567	742	881	968	996
	1	12	49	127	256	427	614	785	909	978	997
	1	10	50	142	297	496	698	857	952	992	999
	1	14	67	187	375	594	788	917	978	997	***
	1	20	86	220	415	629	810	927	982	998	***
	1	15	67	180	355	563	757	896	970	996	***
	0	6	32	101	233	423	636	820	939	990	999
	0	3	16	56	145	296	498	712	883	976	997
	0	3	16	52	130	263	449	659	846	965	996
	0	4	19	61	149	293	485	692	867	971	997
	0	4	23	74	176	337	539	742	897	979	998
	0	7	33	93	203	364	558	747	894	976	997

6522	1	12	48	124	247	412	597	769	900	975	996
	1	13	51	132	263	434	620	787	910	977	996
	1	12	54	146	295	484	678	837	940	988	998
	1	12	56	156	318	520	718	870	958	993	999
	0	10	48	141	301	507	714	873	962	995	***
	0	5	30	104	251	463	689	867	965	996	***
	0	2	16	67	185	384	624	832	954	995	***
	0	1	10	42	123	277	495	727	902	984	999
	0	1	9	37	103	228	415	637	838	964	996
	0	3	15	51	130	267	456	667	851	965	996
	0	5	26	82	192	358	560	756	902	980	998
	0	9	39	111	236	409	605	785	914	982	998

7210	0	4	18	57	134	262	436	634	817	947	992
	0	4	21	66	155	296	481	679	849	959	994
	0	5	27	86	201	374	580	774	914	983	998
	0	7	36	113	255	453	667	844	952	993	999
	0	7	36	115	261	464	681	856	958	995	***
	0	4	26	95	236	446	678	863	964	996	***
	0	2	17	68	189	389	630	837	956	995	***
	0	1	9	39	116	260	469	699	883	978	998
	0	1	5	20	61	151	305	519	750	929	989
	0	1	4	16	50	127	270	480	723	922	989
	0	1	7	26	77	181	352	576	801	955	995
	0	2	13	45	116	242	424	637	832	961	995

5056

Table 9—continued.

7380	0	7	46	150	329	550	754	896	970	995	***
	0	7	46	146	318	532	737	886	966	995	999
	0	6	44	152	344	578	787	920	981	998	***
	0	9	64	211	445	691	870	962	993	999	***
	1	16	91	259	497	728	888	967	994	***	***
	1	16	78	214	419	644	829	940	987	999	***
	0	9	47	141	303	514	725	883	969	997	***
	0	6	33	104	237	428	640	823	941	991	999
	0	5	29	95	221	405	615	802	928	987	999
	0	3	21	84	217	420	647	836	949	992	999
	0	2	18	87	247	486	729	899	976	998	***
	0	3	28	120	305	550	774	918	981	998	***
7785	1	15	64	169	334	536	733	881	964	995	***
	1	17	70	185	362	572	766	903	974	997	***
	1	19	85	225	430	652	832	941	987	999	***
	1	27	120	300	532	747	895	969	994	***	***
	2	31	135	332	575	788	921	980	997	***	***
	1	19	94	261	498	732	895	972	996	***	***
	0	8	46	150	335	571	785	924	984	999	***
	0	4	24	85	212	408	636	831	950	994	***
	0	3	18	68	177	357	584	796	937	992	***
	0	3	19	75	202	407	649	851	963	997	***
	0	4	28	105	265	496	732	901	979	998	***
	0	9	46	143	313	532	744	896	973	997	***
8020	0	3	16	57	143	285	474	678	851	960	994
	0	3	16	53	132	267	452	659	842	960	995
	0	3	18	59	148	297	495	706	878	974	997
	0	4	23	79	195	376	595	796	931	989	999
	0	4	22	78	199	390	618	820	946	993	***
	0	2	13	51	143	307	532	761	923	991	999
	0	1	8	33	99	230	432	670	873	979	998
	0	1	6	26	78	185	358	583	804	954	995
	0	1	5	19	59	144	294	506	740	927	990
	0	1	4	15	50	130	276	492	738	931	991
	0	1	6	24	75	182	359	588	810	955	994
	0	2	12	46	125	265	461	676	857	965	995
9175	3	40	149	332	551	749	888	963	992	999	***
	3	44	151	324	531	725	869	953	989	999	***
	4	55	185	384	603	788	910	972	994	***	***
	2	45	186	419	667	852	952	989	999	***	***
	1	34	165	406	672	865	960	992	999	***	***
	2	40	176	411	666	855	954	990	999	***	***
	2	42	168	380	619	813	931	982	997	***	***
	1	23	103	261	477	694	859	953	990	999	***
	1	14	68	191	381	600	793	920	980	998	***
	2	29	113	267	471	678	841	941	986	998	***
	4	61	203	411	632	810	922	976	995	***	***
	4	59	204	423	651	828	934	981	997	***	***
9330	0	1	12	76	251	526	789	942	992	***	***
	0	2	18	85	247	494	747	915	984	999	***
	0	3	30	121	305	553	783	928	986	999	***
	0	5	44	173	404	668	866	965	995	***	***
	0	6	56	210	467	732	906	979	997	***	***
	0	8	61	210	453	709	889	972	996	***	***
	0	6	44	158	368	622	834	951	992	***	***
	0	2	18	83	238	481	736	912	985	999	***
	0	1	9	47	161	376	647	867	974	999	***
	0	1	11	58	183	402	663	871	973	998	***
	0	2	18	90	259	511	761	922	986	999	***
	0	1	16	94	290	571	819	952	993	***	***
9560	0	9	57	187	403	648	844	953	992	999	***
	0	10	61	190	402	644	840	951	992	***	***
	1	15	84	243	479	719	889	971	996	***	***
	0	14	89	272	534	776	925	984	998	***	***
	0	8	69	246	519	775	928	985	998	***	***
	0	7	62	228	494	754	916	982	998	***	***
	0	8	57	195	422	672	862	961	994	***	***
	0	4	30	109	267	491	721	890	974	998	***
	0	2	13	58	169	364	610	829	955	996	***
	0	3	22	87	233	457	702	886	975	998	***
	0	9	58	191	411	660	855	959	994	***	***
	0	12	75	233	472	716	888	970	996	***	***

8190
 1 19 87 223 415 622 798 915 975 996 ***
 1 17 77 203 387 596 780 907 973 996 ***
 2 29 117 282 495 703 860 950 988 999 ***
 2 39 153 351 584 785 915 976 996 ***
 1 25 116 296 533 753 901 973 996 ***
 0 12 73 227 459 698 873 963 994 ***
 0 8 57 189 406 647 839 947 989 999 ***
 0 4 30 109 268 490 716 883 969 996 ***
 0 1 11 48 146 326 566 794 939 993 ***
 0 2 18 71 195 398 639 842 958 995 ***
 0 13 70 206 415 645 830 940 986 999 ***
 1 24 116 290 511 720 871 954 989 999 ***

8350
 1 14 62 165 327 524 715 864 953 991 999
 1 16 67 172 333 529 720 868 956 993 999
 1 23 96 238 437 649 824 933 998 ***
 1 25 116 297 534 755 902 973 996 ***
 1 18 93 257 491 722 886 968 995 ***
 0 10 60 186 393 631 828 944 989 999 ***
 0 5 35 124 297 530 756 910 980 998 ***
 0 1 13 58 170 364 608 823 951 995 ***
 0 1 6 30 101 250 475 721 904 986 999
 0 2 15 58 159 329 549 765 917 986 999
 0 9 51 150 316 524 725 876 961 994 999
 1 15 73 197 382 591 775 903 970 995 999

8685
 0 2 13 46 117 243 424 634 827 956 994
 0 2 12 42 110 232 411 621 818 951 992
 0 3 19 62 153 302 498 705 873 971 996
 0 6 30 94 217 400 612 803 931 988 999
 0 7 38 116 258 456 670 846 952 993 999
 0 10 49 145 308 518 725 882 967 996 ***
 0 10 52 155 326 542 748 896 973 997 ***
 0 4 24 82 199 381 598 796 930 989 999
 0 0 3 15 54 147 316 549 787 949 994
 0 0 1 6 26 85 215 434 703 920 989
 0 1 6 25 75 177 346 566 788 946 993
 0 3 15 52 130 265 453 664 848 965 996

Table 10
 Monthly model estimates of probabilities that 30 day precipitation sums R are less or equal to 0.05, 0.10, 0.15, 0.30, 0.50, 0.75, 1.00, 1.50, 2.00, 2.50, 3.00 times expected values of R, arranged horizontally. Months and stations are arranged as for Table 9. The numbers for extremely small values of R may be systematically somewhat too small. See section 9 for a possible correction formula.

60	1	4	10	56	177	377	570	829	942	982	995	1835	4	12	25	88	218	404	575	810	923	971	989
	1	4	9	40	136	331	550	860	971	995	999		19	40	64	154	290	454	594	787	893	947	975
	7	18	34	103	236	423	593	825	934	977	992		20	37	57	133	258	422	572	789	905	960	984
	6	15	29	95	225	410	581	816	929	975	991		12	23	38	98	211	381	550	804	931	980	995
	8	16	27	79	187	362	543	816	943	986	997		8	19	33	95	213	388	559	809	932	979	994
	2	5	11	50	153	341	541	830	951	988	998		7	14	25	76	184	361	544	819	945	987	997
	1	2	5	32	125	327	561	879	979	998	***		2	5	10	45	143	331	539	844	963	993	999
	1	3	8	42	145	342	552	843	958	991	998		1	4	9	46	153	350	557	842	956	990	998
	3	9	18	73	199	398	587	840	948	985	996		5	13	26	88	217	408	587	829	939	980	994
	3	10	21	80	209	400	578	820	932	976	992		5	13	25	84	206	386	557	801	921	971	990
	1	5	11	53	165	364	567	845	957	990	998		6	16	30	98	230	420	596	835	942	982	995
	0	2	5	29	115	304	527	853	970	996	***		5	14	29	99	235	423	590	817	926	972	990
165	3	8	16	65	182	374	564	828	944	984	996	2012	5	13	25	87	212	397	572	816	931	977	993
	11	23	38	103	226	408	583	829	942	983	996		10	23	41	115	248	425	584	804	916	966	987
	14	29	47	124	258	434	590	806	915	965	986		15	33	56	142	279	449	596	798	905	957	981
	11	21	34	88	196	366	543	816	945	988	998		11	24	41	113	240	415	576	805	922	971	990
	10	19	31	84	194	369	550	822	947	988	998		9	19	32	91	205	378	550	808	935	982	996
	9	19	32	92	213	393	566	812	930	976	992		5	12	23	74	184	360	539	808	936	982	996
	4	9	17	60	163	348	548	840	959	992	999		2	5	11	47	148	338	543	839	959	992	999
	4	8	16	59	166	354	551	833	952	989	998		3	8	15	59	167	355	549	829	949	987	997
	8	17	30	94	223	412	588	826	936	978	993		9	21	38	108	239	421	587	819	930	976	992
	5	13	24	82	204	389	565	811	928	975	992		5	15	29	97	226	405	568	797	913	965	986
	2	6	13	53	161	362	575	866	971	995	999		6	17	34	111	252	440	604	823	928	973	990
	2	6	13	53	159	352	556	845	960	992	999		3	11	25	92	226	415	585	816	926	972	990
535	3	10	22	85	217	408	583	818	929	974	991	2188	3	9	20	76	201	393	575	824	937	979	994
	6	15	26	83	200	380	556	811	932	978	994		10	22	38	107	234	412	576	807	922	971	990
	15	29	46	114	234	403	564	801	923	974	992		14	28	47	122	252	425	583	805	919	969	989
	10	20	33	90	202	373	545	804	932	980	995		12	24	39	104	224	397	564	810	932	979	994
	9	17	29	80	187	358	538	812	942	986	997		11	23	39	106	229	403	567	803	922	972	990
	5	12	21	69	176	355	543	821	947	987	997		4	9	17	57	152	321	510	810	949	990	999
	2	5	11	46	142	328	537	846	966	995	999		0	1	4	25	110	308	544	870	976	997	***
	2	6	12	53	162	355	556	838	954	989	998		1	4	10	51	165	368	573	848	957	990	998
	3	9	18	66	179	370	566	842	956	990	998		5	12	23	81	208	406	594	848	953	988	997
	5	13	24	82	201	384	561	816	936	981	995		6	15	29	94	220	401	570	806	923	971	990
	3	9	18	66	181	375	575	856	965	994	999		5	11	22	74	190	377	563	827	945	985	996
	2	7	16	69	193	387	571	822	936	979	993		3	9	19	74	198	389	569	818	933	977	993

2780
 7 18 33 105 239 422 586 810 921 968 988
 16 32 54 137 273 445 595 800 908 959 982
 26 52 81 184 330 494 627 805 900 950 974
 23 40 59 131 249 410 563 794 918 972 991
 11 22 36 96 210 384 557 816 940 984 996
 5 13 23 73 181 358 541 815 942 985 997
 4 11 20 66 172 353 544 826 950 989 998
 3 10 21 80 207 398 577 821 934 978 993
 8 18 31 95 222 411 591 837 946 984 996
 5 13 26 87 212 397 570 812 928 974 991
 5 15 28 96 233 426 601 833 938 979 993
 6 17 35 114 253 435 591 803 910 960 983

3490
 5 14 28 96 228 412 579 810 922 970 989
 13 30 51 137 277 451 601 805 911 961 983
 14 33 57 148 290 462 607 803 906 957 980
 15 32 53 136 268 438 587 796 907 960 983
 5 13 25 85 207 390 564 810 928 975 992
 4 10 19 66 174 353 537 812 940 984 996
 3 9 18 71 190 377 560 817 936 980 994
 3 11 24 92 227 417 586 815 925 971 989
 6 15 28 91 218 405 581 825 937 980 994
 4 11 23 85 212 399 571 810 926 973 991
 4 12 25 90 224 414 586 820 930 975 991
 3 10 23 89 224 413 583 814 925 971 990

3845
 4 13 28 99 236 423 589 814 923 970 989
 21 48 80 189 334 495 623 795 891 942 969
 22 46 73 171 312 477 615 802 903 953 978
 13 29 49 130 262 433 585 799 911 963 985
 11 24 40 110 235 411 574 807 923 972 990
 7 16 28 84 195 367 541 804 933 981 995
 6 14 26 84 204 385 558 805 925 974 992
 4 12 25 90 221 407 576 809 923 971 989
 7 18 33 100 228 413 584 821 933 977 993
 5 13 26 89 216 399 568 806 923 972 990
 2 9 20 81 213 409 588 830 939 980 994
 4 15 31 108 247 428 586 802 911 962 984

655
 2 6 13 59 179 379 573 834 945 984 995
 5 12 23 77 192 375 557 820 941 984 996
 11 23 40 110 238 415 579 810 925 973 991
 6 14 26 83 201 384 563 819 939 982 995
 7 14 24 71 173 343 525 808 943 987 997
 3 8 15 55 153 332 530 832 959 993 999
 1 4 8 40 135 329 548 862 973 996 ***
 2 5 11 50 156 350 556 845 960 991 998
 2 7 15 61 175 371 569 844 956 990 998
 3 9 19 72 190 377 561 821 940 982 995
 2 5 12 54 166 365 572 855 963 993 999
 1 4 9 47 154 354 562 846 958 990 998

1010
 1 4 10 57 181 383 575 832 943 982 995
 1 4 12 63 192 396 586 837 945 983 995
 2 7 16 65 181 374 566 832 947 985 996
 3 9 20 79 206 396 572 814 929 975 991
 6 14 25 78 191 371 554 823 945 986 997
 1 5 12 53 161 351 547 826 947 986 997
 0 1 4 28 119 320 549 864 972 995 999
 1 2 6 37 138 343 565 864 969 994 999
 0 2 5 29 118 316 547 870 977 997 ***
 0 2 5 34 133 334 554 855 966 993 999
 0 2 6 40 147 346 552 836 952 988 997
 1 4 11 62 190 390 577 826 937 979 993

1252
 6 15 29 92 218 399 567 804 921 971 990
 11 22 36 96 213 388 561 815 938 982 996
 27 48 72 158 289 453 595 796 904 956 981
 19 38 60 139 265 427 573 788 906 961 985
 13 25 41 103 218 386 551 800 926 976 993
 8 17 30 87 200 376 552 812 937 982 995
 2 6 12 49 149 336 542 844 963 993 999
 4 11 21 76 195 382 564 820 937 981 994
 6 14 24 77 191 375 564 837 955 990 998
 6 14 26 81 195 371 546 804 931 979 994
 7 18 35 113 255 442 603 817 922 967 987
 4 11 22 82 210 402 581 824 935 978 993

Table 10—continued

3922	2	7	17	74	204	402	585	830	940	980	994
	11	28	52	144	290	465	611	808	909	958	981
	11	27	47	128	266	445	601	814	920	968	988
	8	18	30	88	204	382	560	822	943	985	997
	10	21	34	93	208	384	558	814	937	982	995
	6	13	23	71	176	352	537	816	945	987	998
	4	12	24	82	202	384	557	804	925	974	991
	2	6	13	58	171	362	552	820	940	982	995
	3	7	15	58	166	353	548	829	950	988	997
	2	6	12	53	160	350	548	831	951	988	997
	1	5	12	59	180	383	582	846	954	988	997
	1	5	14	69	198	395	579	825	936	979	993
4289	4	11	22	81	208	397	574	816	930	976	992
	12	28	48	131	268	445	600	811	918	966	987
	31	67	104	220	363	511	628	786	877	929	959
	6	15	29	93	221	408	582	823	935	978	993
	7	18	33	101	230	410	575	806	922	971	990
	3	9	18	66	175	359	551	831	953	989	998
	1	4	9	47	154	352	559	846	959	991	998
	1	4	9	47	155	352	556	837	952	988	997
	2	5	12	54	165	359	558	835	951	988	997
	6	14	26	80	196	379	561	823	942	984	996
	1	5	11	49	153	349	558	852	964	993	999
	3	10	21	83	216	410	588	827	936	978	993
4448	5	14	26	86	212	401	578	822	936	979	994
	7	18	34	106	242	429	597	823	930	975	991
	14	32	55	144	284	455	600	798	903	955	979
	4	9	18	67	179	368	561	836	953	989	998
	7	18	33	99	227	410	580	817	931	976	992
	2	5	11	43	131	305	510	832	964	995	999
	1	4	10	45	145	338	547	844	961	992	999
	1	3	7	38	137	335	548	846	961	992	999
	2	5	10	46	147	345	560	861	970	995	999
	3	8	17	65	182	377	569	835	949	986	997
	0	2	4	21	91	275	522	886	987	999	***
	2	6	13	57	174	378	581	854	960	991	998
5147	6	16	30	95	223	406	574	809	924	972	990
	13	31	54	141	281	455	605	810	915	964	985
	24	52	84	192	336	493	620	791	887	939	968
	9	22	39	112	243	423	584	809	921	969	989
	7	16	30	92	216	397	568	811	928	975	992
	5	11	19	58	151	322	519	832	964	995	***
	1	3	8	40	137	331	544	846	962	993	999
	1	2	6	30	115	305	534	865	976	997	***
	1	3	7	37	131	327	549	862	972	996	999
	3	10	20	79	210	408	592	838	945	983	995
	2	5	10	43	137	324	537	850	968	995	999
	7	18	33	101	233	420	592	825	935	978	993
5711	5	13	24	83	206	391	566	813	930	976	993
	8	21	39	114	249	430	591	813	923	971	989
	11	28	51	142	285	460	606	805	908	958	981
	3	11	22	84	215	408	586	826	936	978	993
	7	17	32	102	235	418	584	811	923	971	989
	2	6	13	52	154	337	536	828	952	989	998
	3	10	21	82	212	404	581	822	933	977	992
	2	5	11	50	155	348	550	838	955	990	998
	1	3	7	38	136	335	554	856	967	994	999
	2	5	10	48	156	359	572	861	967	994	999
	0	2	4	25	104	296	533	874	980	998	***
	2	7	16	66	187	382	571	830	944	984	996
5896	9	20	36	105	233	412	577	807	921	970	989
	11	27	49	137	279	455	605	809	913	962	984
	15	35	60	153	295	462	602	794	897	950	976
	5	15	31	107	250	439	602	818	922	968	987
	7	15	26	81	198	382	562	818	937	981	995
	5	13	26	86	210	394	567	811	928	975	992
	1	4	10	50	160	357	558	835	951	987	997
	1	4	9	40	133	322	539	852	968	995	999
	1	4	8	39	132	327	549	866	975	997	***
	4	10	20	72	194	391	584	845	953	988	997
	2	5	11	49	150	337	538	829	952	989	998
	7	17	31	98	228	412	581	814	927	973	991

4480
 4 9 18 67 183 375 567 835 949 987 997
 5 14 27 93 224 415 591 829 938 980 994
 7 19 37 114 251 430 586 800 909 961 983
 3 8 17 64 177 368 562 835 951 988 997
 3 9 18 67 181 369 557 823 942 983 995
 3 8 15 54 149 327 528 838 964 995 999
 1 4 8 38 128 314 531 850 969 995 999
 1 2 5 33 128 328 549 854 966 994 999
 1 2 6 32 123 322 550 866 973 996 ***
 3 8 15 56 163 356 560 847 961 992 999
 1 2 4 22 88 265 511 891 990 *** ***
 3 7 14 55 162 359 567 857 966 994 999

6522
 10 23 41 115 248 426 586 808 919 968 988
 9 20 40 117 252 433 594 816 925 972 990
 13 30 51 136 271 441 589 794 904 957 981
 5 14 28 97 233 421 590 819 928 974 991
 4 10 20 71 188 375 561 824 942 983 996
 2 5 12 54 166 360 558 833 949 987 997
 1 3 7 40 142 341 556 852 964 993 999
 1 3 8 42 146 345 556 848 961 992 998
 1 4 9 42 139 330 543 848 965 994 999
 3 10 20 77 203 401 589 844 951 987 997
 3 10 21 77 199 384 561 809 928 976 992
 13 31 53 142 281 454 602 804 909 960 983

7210
 3 8 18 72 194 387 572 827 941 982 995
 3 9 18 68 186 377 565 828 944 984 996
 11 28 51 140 282 455 602 801 906 957 981
 4 12 24 83 206 392 569 818 934 979 994
 2 6 12 50 149 333 534 830 955 990 998
 2 6 12 50 153 342 546 840 959 992 999
 1 4 9 43 142 335 548 850 965 994 999
 1 2 5 29 110 297 526 864 977 997 ***
 1 4 9 45 150 347 554 843 958 991 998
 1 3 8 41 144 347 566 863 969 994 999
 1 4 10 52 165 361 557 828 945 984 996
 5 13 25 89 221 412 587 825 935 978 993

7380
 3 10 22 87 220 410 583 817 929 974 991
 4 14 30 107 247 429 586 801 910 961 983
 6 23 47 146 298 473 614 799 897 948 973
 1 3 7 42 151 353 559 839 952 987 997
 3 8 15 59 170 362 561 841 956 990 998
 2 6 11 43 133 313 524 845 968 996 ***
 2 5 11 48 148 338 547 847 963 993 999
 2 7 14 53 154 343 550 853 969 995 999
 0 2 5 33 131 336 556 854 963 992 999
 1 4 9 51 167 373 580 854 960 990 998
 0 2 7 47 161 359 555 823 940 981 995
 1 5 14 72 208 410 592 830 936 978 992

4645
 9 21 37 109 239 416 575 798 912 964 986
 20 40 63 151 287 456 603 805 912 962 984
 36 73 112 229 371 517 632 787 876 928 958
 11 24 41 114 245 424 585 810 922 970 989
 11 26 45 124 259 434 589 801 911 962 984
 7 15 25 74 178 350 534 817 948 989 998
 2 5 12 56 170 364 559 829 946 985 996
 1 4 8 36 124 311 532 856 972 996 ***
 2 6 13 58 174 374 574 845 955 989 998
 6 16 31 98 228 414 584 817 929 975 991
 2 7 15 62 178 373 567 835 949 987 997
 7 19 36 115 254 437 596 811 918 966 987

5035
 5 13 26 87 212 397 570 812 929 975 992
 13 32 56 150 297 475 619 815 915 962 983
 18 43 72 175 319 481 613 792 890 943 971
 8 22 41 125 267 448 603 811 916 964 985
 9 25 46 133 274 449 598 801 907 958 982
 3 9 18 64 174 359 549 824 946 986 997
 3 9 20 80 209 399 575 816 929 975 991
 1 4 10 45 147 342 552 847 961 992 999
 2 5 11 49 155 352 561 851 963 993 999
 2 8 17 70 194 392 581 838 948 985 996
 1 3 7 36 125 313 534 856 972 996 ***
 6 16 32 105 243 431 597 821 928 973 990

Table 10—continued

7785	8	21	39	117	252	428	583	798	909	961	984
	15	36	61	156	298	467	606	797	899	951	976
	33	72	112	233	377	523	636	788	876	926	956
	22	53	89	203	351	506	627	789	881	932	961
	8	17	31	90	207	385	559	813	935	981	995
	4	9	17	60	164	351	552	843	961	993	999
	2	4	10	43	140	333	550	858	970	995	999
	1	4	8	37	127	317	539	862	974	996	***
	1	3	7	35	125	319	546	867	975	997	***
	2	6	14	65	190	388	576	828	940	981	994
	3	9	19	78	203	390	564	806	923	971	990
	15	39	68	173	321	487	620	798	895	946	972
8020	4	11	22	78	199	388	572	829	944	984	996
	3	8	16	64	179	371	562	830	946	985	996
	6	15	28	91	217	404	579	822	935	978	993
	2	6	14	63	184	383	576	837	948	985	996
	3	9	18	62	166	345	533	815	944	986	997
	2	7	15	69	196	392	576	826	938	980	994
	5	15	31	108	247	430	590	806	915	964	986
	4	11	23	85	213	400	574	815	929	975	992
	0	1	4	24	106	303	543	877	980	998	***
	1	3	8	39	141	347	571	873	974	996	***
	1	4	10	49	154	344	543	826	948	987	997
	4	10	21	74	193	383	570	832	947	986	997
8190	8	22	42	121	256	430	583	796	908	960	984
	7	18	35	109	243	424	585	806	918	967	987
	18	40	66	161	301	466	604	793	896	949	975
	15	35	60	153	295	464	605	796	899	950	976
	7	18	35	110	243	420	575	790	903	956	981
	2	6	13	58	173	372	573	847	957	990	998
	2	6	14	60	174	367	560	830	946	985	996
	1	3	7	40	141	340	555	852	964	993	999
	0	2	6	40	146	346	553	837	953	988	997
	3	9	20	80	213	410	591	832	940	980	994
	5	14	29	99	230	410	572	797	911	963	985
	12	29	50	135	271	445	596	805	913	963	985
9330	1	8	20	88	229	422	589	811	919	966	986
	6	21	43	135	282	460	606	802	904	955	979
	2	7	17	77	208	403	581	821	932	976	992
	2	6	13	55	162	351	549	834	953	989	998
	4	14	30	104	240	421	580	797	908	960	983
	2	5	11	48	154	355	570	866	971	995	999
	2	7	16	65	181	371	559	822	940	982	995
	1	3	7	36	131	325	542	851	966	994	999
	0	1	4	27	114	309	535	854	968	995	999
	1	4	9	50	164	367	571	845	955	989	997
	0	2	4	27	113	311	547	879	981	998	***
	0	2	6	40	150	359	572	853	960	990	998
9560	2	5	11	49	152	340	543	837	957	991	998
	3	10	21	81	209	401	578	820	932	976	992
	4	11	22	74	191	378	563	825	942	984	996
	5	13	26	89	217	405	580	821	933	977	993
	2	6	14	57	168	357	550	826	946	986	997
	1	4	10	47	152	349	560	854	965	994	999
	4	11	22	82	210	402	581	826	937	980	994
	1	3	8	37	128	315	533	851	968	995	999
	0	2	5	31	123	321	544	855	967	994	999
	2	5	10	46	148	349	571	876	978	997	***
	1	5	11	51	161	358	559	836	950	987	997
	2	8	17	66	184	376	568	835	950	987	997

8350
 14 32 54 137 270 439 587 794 905 959 983
 17 18 35 107 240 420 581 805 918 967 988
 17 38 64 158 299 468 609 801 903 954 979
 18 41 68 164 304 468 605 794 897 949 976
 6 15 29 94 219 399 565 800 918 968 988
 3 8 16 65 184 380 573 839 951 987 997
 2 5 11 49 151 339 538 827 950 988 997
 1 3 8 46 158 358 561 836 950 986 997
 0 1 5 33 130 327 542 840 957 990 998
 4 11 24 88 223 419 596 833 939 980 994
 4 12 23 77 193 373 552 812 935 981 995
 8 19 35 104 232 410 576 812 928 975 992

8685
 1 5 11 48 151 344 550 843 960 992 999
 1 3 7 38 134 332 551 858 969 995 999
 2 6 12 52 159 356 565 854 964 993 999
 4 10 19 71 189 382 573 837 950 987 997
 2 7 14 57 166 355 550 828 948 986 997
 4 11 20 68 178 368 566 848 962 993 999
 6 15 28 87 206 387 561 813 933 979 994
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 0 2 5 32 133 344 573 873 973 995 999
 0 2 4 30 122 322 546 857 968 994 999
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 10 22 39 110 238 412 571 796 913 965 986
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 10 22 37 101 220 396 568 821 942 985 997
 21 38 58 136 265 436 592 811 922 971 990

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